

**A SUMMARY OF ONE YEAR OF OBSERVATIONS
OF 6300A AIRGLOW AT HALEAKALA**

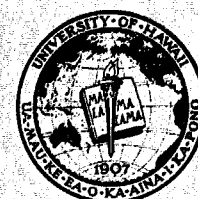
By
F. E. ROACH, W. R. STEIGER, and W. E. BROWN

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TECHNICAL REPORT

Prepared for
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UNIVERSITY OF HAWAII



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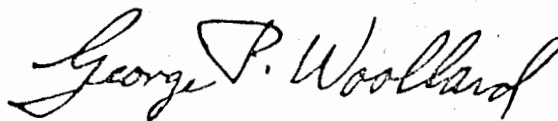
Technical Report

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UNDER NASA GRANT NsG-135-61

Approved by Director

A handwritten signature in cursive script, reading "George P. Woolland". The signature is written in dark ink and is positioned below the "Approved by Director" text.

Date: 2 October 1964

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A SUMMARY OF ONE YEAR OF OBSERVATIONS
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I. Introduction

The atomic oxygen transitions $^3P_{2,1} - ^1D_2$ or $[OI]_{21}$ yield two prominent radiations in the night airglow at wave lengths of 6300.3A and 6363.8A. The excitation potential is 1.96 e.v. The "forbidden" nature of the transitions is apparent from the mean life of the 1D_2 state, 110 seconds.

The radiation 5577.3A is also due to a forbidden transition $^1D_2 - ^1S_0$ or $[OI]_{32}$. The energy level diagram (figure 1) illustrates that the two radiations are related through their common energy level, 1D_2 . After the emission of the green, 5577A, line there are two possibilities: the atom may (after ~ 110 seconds) emit the red (6300A, 6364A) lines or it may be deactivated by collision.

In most airglow (and auroral) manifestations the green and red emissions seem to have little in common. Their diurnal variations are in general so dissimilar (figures 2 and 3) as to imply entirely different excitation mechanisms. An oversimplified conclusion from this lack of co-variance of the two emissions is that the green line is produced at lower heights (~ 100 km) where the collisional deexcitation of the 1D_2 level discourages the emission of the 6300A* line; and that the red line occurs predominantly at greater heights (which we may designate generally as the F region) where collisions are infrequent. A corollary conclusion from the lack of co-variance is that excitation mechanisms occur in the F region that are more effective at 1.96 e.v. than at 4.17 e.v.

*Henceforth we shall refer to 6300A emission with the understanding that we include also 6364A.

Direct measurements of the height of the emissions by rocket explorations are not inconsistent with the above interpretation. A number of measurements in mid-latitudes have placed the green line emission layer in the 100-kilometer region [Packer, 1961; Huruata, 1962]. To date there has not been an isolation of the 6300A emitting layer but in one rocket ascent it was established that it was greater than 200 kilometers [Zipf and Heath, 1962].

Some general evidence favoring the F region for the emission of the red line may be mentioned. In the case of visual auroras of Type A the upper red border, due chiefly to 6300A, is higher than the lower green part of the aurora. Another 6300A phenomenon, the mid-latitude M-arc, has been found by triangulation to be at a height of about 400 kilometers [Roach and Roach, 1963]. Also it should be mentioned that tropical 6300A arcs have been shown by Barbier, et al [1961] to occur in the 275 kilometer region.

If one considers the world-wide aspects of the 6300A - 5577A problem, one of the impressive facts is the very wide spread of their relative and absolute intensities for the several different phenomena. In Table 1 we show that the intensity ratio, $\frac{Q(6300)}{Q(5577)}$, covers a range from the order of 80 or greater for M-arcs to $\frac{1}{5}$ for the quiescent airglow. The conclusion that is forced on us from this spread is that we must be dealing not only with a difference in the preferred emission heights but also with a variety of excitation conditions.

That we are concerned with many different kinds of excitation in the case of 6300A is indicated by the large number of different phenomena that have been isolated. Barbier [1958, 1961] has identified the following at Haute Provence (HP) and at Tamanrasset (T):

- The polar aurora (HP)
- The twilight phenomenon (HP, T)
- The western sheet (HP)
- The subpolar sheet (HP, T)
- The tropical arc (T)
- The mid-latitude arc (HP)
- The "para" twilight phenomenon (T)

The purpose of this paper is to summarize more than a year of observational data on the 6300A emission made at the tropical station at Haleakala (geographic latitude N 20° 42' 43", geographic longitude W 156° 15' 47", magnetic latitude 20° 50' N, magnetic longitude 88° 27' W, altitude 10,012 feet or 3052 meters, A.S.L. on the island of Maui, Hawaii). We propose to give a general description of the observations with particular emphasis on the marked intensity enhancements which have frequently occurred.

II. The Observational Material

The airglow observing program at Haleakala involves two photometers:

(1) a zenith photometer that records the night sky intensity successively in a 5-minute sequence through filters centered on wave lengths 5577, 6300, 5893, and 5300A [Purdy, Megill, and Roach, 1961] and (2) an alt-azimuth mounted photometer that scans the sky in a series of almucantars (zenith distances 80°, 75°, 70°, 60°, 40° and zenith) in a 15-minute sequence in wave lengths 5577, 6300, and 5893A [Roach, Megill, Rees, and Marovich, 1958]. The zenith photometer has a built-in standard light making absolute calibrations possible. The alt-azimuth photometer includes a modified Lyot birefringent quartz filter that isolates the monochromatic airglow emissions from the more or less continuous background of integrated starlight and zodiacal light.

During the 1961-62 observing period the recording for both photometers was on pen-and ink-strip charts. The large quantity of observational material covered by this report was made tractable by putting the measurements into digital form on punched cards which then could be processed by electronic computers available to us.

The method used in converting the readings into absolute units involves the calibration of the built-in standard light of the zenith photometer which has been a major preoccupation of the Fritz Peak Observatory. The reader is referred to a paper by Roach and Smith [1964] for a discussion of the problems involved in isolating the airglow from the background light.

The correction of the readings to what is called the local zenith intensity (LZI) outside the atmosphere has a long history starting with a paper by Roach and Pettit [1951]. In summary the procedure consists in: (1) subtracting out the scattered light from the lower atmosphere, (2) correcting for the extinction of the lower atmosphere, (3) correcting for the Van Rhijn increase of the line-

of-sight brightness toward the horizon, and (4) multiplying by the absolute calibration coefficient to put the results in rayleighs. The processed data are printed out and in some cases put on a new deck of punched cards.

In Table 2 we show a log of the 149 nights included in the present report. In figures 3 and 4 we have plotted representative zenith intensities of 6300A versus time based on the zenith photometer observations. We note that there seem to be two classes of nocturnal variation: (1) what might be called the "classical" variation with the high post-twilight intensity decreasing rapidly (like a negative exponential curve) to a low value and (2) the classical curve interrupted by one or more dramatic increases in intensity extending over a period of about two hours. The classical variation is characteristic of the experience of observers in mid-latitudes (see figure 2). The abrupt increases illustrated in figure 4 are characteristic of the tropics and have been observed by Barbier [1964], A. and D. Delsemme [1960], Barbier, Roach and Steiger [1962], Carman and Gibson-Wilde [1964], Saito [1962], Davis and Smith [1964], Silverman and Casaverde [1961]. Observations over the sky indicate that the photometric increase covers a significant area with localized spotty regions sometimes arrayed in an arc-like orientation.

If we assume that the effective height of the 6300A emission is 250 kilometers, then our outer observational circle (80° zenith distance) has a radius of 1000 kilometers and includes a region of the eastern Pacific centered on the Hawaiian Islands, extending about half way to the equator on the south. The geographic and geomagnetic latitudes are practically identical, the geomagnetic equator crossing the geographic equator to the south of the Hawaiian Islands. The

geomagnetic parallels are inclined about 11° E to the geographic parallels.

The all-sky observations from the birefringent photometer have been digitized on punched cards by a two man-year measurement program. Values of LZI are printed out in two ways: (1) a simple tabulation of the results as shown in Table 3 and (2) a print-out on a polar coordinate circle map with each value of LZI printed out in its position with respect to a geographical circle centered on the observer (see figure 5). An interpolation program is also included to indicate the positions on the circle plot of isophotes corresponding to sequentially selected values of the isophotes in rayleighs (see figure 6).

From this large mass of data we are reproducing in the present report in Table 4 the zenith values of the intensity of 6300A in rayleighs for 15-minute intervals.

III. The Correlation Between Q(6300) and the Ionospheric Parameters

Some time ago St. Amand [1955] reported a possible relationship between the 6300A mid-latitude airglow and the ionization in the F region. Barbier [1959] made a somewhat similar study of his observations at the Haute Provence Observatory. He assumed that the 6300A line is produced by a charge exchange followed by a dissociative recombination, as follows:



in which rate of production of the excited oxygen atoms, O^* , which eventually emit 6300A, is proportional to the product of the O_2 number density, $[O_2]$, and electron number density, n_e . Hence the number of emissions per second in a vertical column 1 cm^2 in cross-section will be

$$Q = k \int_h^\infty [O_2] n_e dh. \quad (3)$$

By assuming mean values of these quantities in the F region, Barbier infers a semi-empirical relationship between the airglow emission, Q , and the ionospheric parameters f_oF2 (critical frequency of the F2 region) and $h'F$ (virtual height of the F2 region), as follows:

$$Q = k [O_2]_0 (f_oF2)^2 e^{-\frac{(h'F-h_0)}{H}} \quad (4)$$

where h_0 is a reference height (~ 275 km) at which $[O_2]_0$ is taken, H is the O_2 scale height, and K is an empirical constant.

On the basis of a more detailed analysis of the reaction rates of (1) and (2), and considering also the reactions



Peterson [1964] has computed the rate of emission of 6300A as a function of height. Assuming a Chapman-layer distribution for the electrons, he has carried out the integral over height and arrived at an expression for the emission rate, Q, essentially identical to equation 4, in which K involves the scale height, H, and also the effective rate coefficient of reaction (6).

Barbier [1959] found that the correlation between airglow observations in 6300A at Haute Provence and the computed values from his semi-empirical formula was only fair, and in some months very poor. Such was also the case at Maruyama as reported by Huruhata, et al [1959].

In the tropics, on the other hand, an equation of the form

$$Q(6300) = A + B (f_0 F_2)^2 e^{-\left(\frac{h'F - h_0}{H}\right)} \quad (7)$$

can be adjusted to fit the observations with a very high correlation on individual nights. Even on those nights when there are sporadic and large changes in intensity, the formula works very well. An example of such a night is shown in figure 7 where the observed and calculated intensities are shown. It has been found, however, that the constants A and B vary from one observing period to the next and probably also from night to night. Hence, the larger the block of data taken, the poorer the correlation will appear to be. In figure 8 we show the observed Q(6300) versus the calculated function $E = (f_0 F_2)^2 e^{-\left(\frac{h'F - 200}{40}\right)}$ for the 93 nights listed in table 5. The regression line gives $A = 26.8$ and $B = 4.2$ with a correlation coefficient of 0.74.*

*The statistical method used here assumes that E and Q are equally uncertain, both being derived from experimental procedures of comparable accuracy. Thus, the regression line was found which minimized the square of the deviation of the experimental points from the regression line as measured perpendicular to the line. In general this line bisects the regression lines of Q on E and E on Q.

The correlation between the observed Q(6300) from Haleakala and the quantity A + B E has been calculated for various values of H. These results are shown in Table 6. It is evident that the correlations are in general quite high and that the best values result for a scale height in the range of 40 to 50 km.

On the night of Sept. 11/12, 1961, there were two very dramatic enhancements in 6300A. Although, under normal or quiescent conditions, the 5577A emission is quite unrelated to the 6300A emission, on this night enhancements in 5577A were observed in close coincidence with those in 6300A, as seen in figure 4. This suggests that a similar or related mechanism was responsible for the enhancements in both wave lengths.

As a working hypothesis, we suggest that the observed 5577A is due to emissions at two different heights, as illustrated in figure 9. The major component has its origin in the vicinity of 100 km and is quite unrelated to 6300A which has only a single component in the vicinity of 300 km. A second and usually minor component of 5577A appears in the vicinity of 300 km and is related to the 6300A emissions, perhaps through reactions (1) and (2) which are known to be capable of producing both 5577A and 6300A. Under conditions when the 5577A is rather strong (several hundred Rayleighs) in the major component and 6300A is relatively quiescent, then the minor component of 5577A which is correlated with 6300A will be rather insignificant and difficult to detect. However, when the "major" component of 5577A is unusually weak (100R or less) and 6300A is active with large enhancements of several hundred Rayleighs, then the "minor" component of 5577A correlated with the 6300A may become quite significant. This appears to be the case with the example illustrated, for the background or uncorrelated component of 5577A is unusually weak whereas 6300A shows unusually large enhancements. Similar examples

can be found on other nights and a very good example is described by Saito [1962] at a latitude of 20° N on November 13, 1960. (see figure 10)

In view of the relationship between the ionospheric parameters and the observed airglow intensity, it appears that it may be feasible to extend our knowledge of the ionosphere by way of the all-sky airglow observations. Using the zenith observations of both airglow and ionosphere to determine the empirical constants A and B in the Barbier formula (7), one can then apply this formula to other parts of the sky. Unfortunately, there are two unknown ionospheric parameters and so the information is not sufficient. Because f_0F2 does not vary greatly over the visible sky at a given time, one could make an approximate solution for $h'F$ by assuming a constant f_0F2 over the sky equal to the zenith value. The result of such a calculation is shown in figure 11. In this way one could construct a "topographic" map of the ionosphere depicting $h'F$ over the entire visible sky.

In the case of the September 11/12, 1961, data, the Barbier formula allows a rather good fit to the 6300A observations but not the 5577A. It was found that a relationship

$$Q(5577) = 70 + 200 e^{-\left(\frac{h'F-200}{40}\right)} \quad (8)$$

worked rather well for the 5577A observations, as shown in figure 12. Just why this should be is not clear. The rate of production of O^{**} in reactions (1) and (2) and hence the rate of emission of 5577A, is proportional to n_e at heights where n_e is large in comparison to $[O_2]$ and is proportional to n_e^2 at lower heights where n_e is small in comparison to $[O_2]$ [Ratcliffe, 1956]. At no time should it be independent of n_e , as seems to be the case in this example.

In cases where Q(5577) is also found to be related to the ionospheric parameters, such as the night of September 11/12, 1961, mentioned earlier, it is possible to solve for both $h'F$ and f_oF2 from the airglow observations. One must assume, of course, that the two empirical formulas for 5577A and 6300A hold over the entire sky. This has been done for the night of September 11/12 and the results are shown in figure 13(a) and (b). In (a) are shown lines of equal height ($h'F$) and in (b) are shown lines of equal critical frequency (f_oF2). To the extent that such an analysis is valid, the results are quite dramatic and should be of considerable interest to the ionospheric physicist. One can make the observation that the variations in the height $h'F$ over the field of view of radius approximately 1000 km are of the order of 5 to 10% of the horizontal dimensions, which is far greater than one finds on the surface of the earth where the variations are more like a fraction of 1%.

It is of interest to take note of the work of J. W. Wright [1959] who has produced vertical profiles of electron density along a section through the 75th meridian. These profiles show changes in height of the ionosphere along a North-South line. Our maps give somewhat similar information but in two dimensions. The decrease in height observed in the vicinity of 15° latitude is consistent with our observations. In figure 14 is a direct comparison of Wright's contours and ours.

IV. The Relationship Between Tropical Q(6300) and the Geomagnetic Field

Both the photometric (6300A) and the F-region ionospheric (f_0F_2 , $h'F_2$) phenomena show a synoptic dependence on the geomagnetic field. Each has a region of "singularity" about $\pm 15^\circ$ from the magnetic (dip) equator. Direct evidence for the 15° preference in the case of 6300A has been obtained by Barbier, Weill, and Fafiotte [1961] and Barbier, Weill, Daguillon, and Marsan [1961] in two round-trip flights from France to South Africa. Two of Barbier's plots of Q(6300) versus dip latitude are shown in figures 15 and 16.

Similar observations on slow-moving ships are more difficult to interpret since seasonal and/or daily changes may mask the geographical term. The zenith observations from the research ship Eltanin as reported by Davis and Smith [1964] show a definite enhancement of 6300A at $\pm 15^\circ$ magnetic latitude (figure 17).

The Japanese ship Soya has made several voyages between Japan and Antarctica with photometric equipment aboard. In figure 18 taken from Huruhata [1963] we show composite results illustrating the latitude changes of 6300A.

The literature on the F-region behavior in the tropics is extensive. In 1954, Appleton called attention to what he called a geographic anomaly in the distribution of f_0F_2 . In figure 19 we show his plots of f_0F_2 versus magnetic latitude during the equinox for 0900, 1200, and 2100 local time. The "anomaly" shows up at the latter two times as a "bite out" or secondary minimum near the magnetic equator and maxima some 10° to 15° from the equator.

The effective height of the F layer and its change with latitude (magnetic) is shown in figure 14 taken from J. W. Wright [1959]. The iso-height lines clearly minimize about 15° from the magnetic equator.

We thus have a consistent picture of two phenomena, photometric and ionospheric, which correlate with each other via the Barbier formula and which both show a singular region approximately 15° from the magnetic equator. Later we shall review some of the attempts to rationalize the behavior of the F region in the tropics but here wish to return to the matter of its magnetic control.

A natural question arises as to whether there is any observable relationship between the tropical 6300A emission and planetary magnetic activity. In figure 20 we show a plot of nightly means of the zenith Q(6300) at Haleakala, including all of the data for a year, versus the planetary magnetic index Kp. There does not appear to be any evident relationship indicated by this scatter of points, but it is possible that the seasonal variation in Q(6300) (see Section VI) may be masking the effect. Thus, the seasonal variation was removed from the data and replotted in figure 21. There is no evident improvement as far as showing any relationship between the Q(6300) and Kp. It was then thought that since September, October, and November are the most active months in Q(6300), that these should be considered by themselves. Thus, in figure 22 we show only the Q(6300) data for those months versus Kp. There is perhaps a slight hint of some relationship evident, but we do not consider it significant. Removing the seasonal variation does not change the situation.

Various other things were tried such as using Q(6300) data from 80°S and 80°N along the meridian, but none of these showed any relationship to Kp.

We must conclude, then, that with this set of data and insofar as we have analyzed it, there does not appear to be any significant relationship between Q(6300) at Haleakala and the planetary magnetic index Kp. This result is perhaps somewhat surprising in view of the preceding statements about the magnetic control of the 6300A emissions.

V. The Alignment of Q(6300) Isophotes

The isophote maps can be roughly classified into two types: (1) those that show a "spotty" character with at least some of the isophotes making closed loops; and (2) those that show a somewhat "aligned" character with the isophotes roughly linear and parallel to each other.

The spotty character is often associated with the localized enhancements that have been discussed in some detail elsewhere in this report. The aligned character is almost always associated with morning and evening twilight effects, but also during the middle of the night there often appears to be a definite alignment.

It was felt that the directions of these alignments might not be purely random and that a determination of the preferred orientation, if any, would be of interest. For this study, the night was divided into three parts: (1) the "evening" period going from twilight to 1955; (2) the "midnight" period from 2155 to 0255; and (3) the "morning" period from 0410 into twilight. For each 15-minute interval during these periods for which an isophote map was available a determination was made as to whether the map was spotty or aligned. If aligned, the direction of alignment in the vicinity of the zenith was determined the direction being defined such that the brighter sky was to the left.

The analysis of 4,617 6300A isophote maps from July 1961 to July 1962 showed that 1,602 possessed an aligned character, and of these 188 were during the evening period, 728 were during the midnight period, and 180 were during the morning period. Figure 23 illustrates the frequency of occurrence of alignments during the year for the midnight period. Evidently they occur somewhat more frequently during the fall months than during the rest of the year, the period, also, when the average Q(6300) reaches a seasonal maximum. (see Section VI)

In figure 24 is plotted the azimuth of each evening alignment during the year. The solid curve on the graph represents the azimuth of the setting sun plus 90° . It is seen that the general trend of the alignments is to follow the solid curve, indicating that the increasing gradient of the evening airglow points in the general direction of the setting sun. This result is not at all surprising; rather, it is surprising that it does not point exactly in the direction of the setting sun, for it is seen that in the summer months the gradient points somewhat to the north of the setting sun, and somewhat to the south of the setting sun in the winter months.

The results for the morning period are plotted similarly in figure 25. Again, the same general conclusion follows: that the increasing gradient of the airglow points in the general direction of the rising sun, but in this case there is a rather consistent tendency for the direction to be somewhat north of the rising sun.

The midnight period maps showed alignments in all directions but a very high preponderance of alignments with an azimuth in the range of 281° to 290° , as illustrated in figure 26, which shows the number of alignments measured in each 10° interval of azimuth. The same data ^{are} ~~is~~ presented in figure 27 in the form of a polar plot. Shown also in figure 27 is the azimuth of the horizontal component of the geomagnetic field, 11° E of N.

One sees at once that the predominant direction of the isophotes is perpendicular to the direction of the geomagnetic field, with the gradient in the direction of the field, increasing from north to south. It seems highly unlikely that this is merely a coincidence and, in fact, is consistent with the concept of geomagnetic control of the 6300A emission as well as of the F region of the ionosphere as discussed in some detail in the previous section.

VI. The Seasonal Variation of Tropical Q(6300)

The airglow is a very complex phenomenon, as was pointed out in the introduction. It is almost certain that no single mechanism can account for all its aspects. Contributing to the total phenomenon there must be numerous factors such as the geomagnetic field, the equatorial electrojet, photochemical reactions, electric fields, collisions, solar radiation, etc. The search for seasonal variations in selected aspects of the airglow is an attempt to determine if there is any solar control and, if so, by what mechanism does it operate.

One of the aspects selected for study has been the position of the maximum intensity of Q(6300) along the meridian immediately after the end of twilight. If the latitude of the sun does have an influence on the distribution of the airglow, one would expect it to be most pronounced in the early evening. The observations do, in fact, show that there is a seasonal variation in the position of this maximum, moving southward in winter and northward in summer, in phase with the varying declination of the sun, as shown in figure 28.

These results were obtained from the analysis of data from 149 nights over a period of 13 months. The position of the evening maximum was measured on each night and these were averaged for each month. The results are shown in Table 7 in terms of the latitude of the maximum as well as the actual distance north (+) or south (-) of the observing site along the surface of the earth. The column headed "L" gives the value of McIlwain's [1961] "sheet parameter"* at a height of 300 km corresponding to the latitude of the maximum.

This seasonal variation is illustrated in another way in figure 29 where the relative intensity of Q(6300) just after twilight along the meridian has

*In a perfect dipole field, L is the geocentric distance to the intersection of the pertinent line of force and the equatorial plane. In the actual field of the earth, L is approximately the same quantity as McIlwain uses. Points having a given value of L define a sheet on which a trapped particle is constrained to move.

been plotted for a number of nights during the year. The movement of the peak towards the south in the winter and towards the north in the summer is quite clear.

Another seasonal variation can be found by plotting the 3-month running means of the zenith intensity of 6300A during the hours 2200 to 0300 HST. The result is shown in figure 30, in which a pronounced maximum appears in September and a minimum in the months of January, February, and March. This is quite the reverse of the result obtained by A. and D. Delsenme [1960] who found a distinct maximum in March and a minimum in September at the southern hemisphere station (4° S Geomagnetic) at Lwiro. This suggests that the maxima and minima in both hemispheres occur at the equinoxes but out of phase with each other.

Still another device for illustrating a seasonal variation is shown in figures 31, 32, and 33. Here are arranged the mean hourly intensities for each hour of the night, for each month of the reporting period. The array becomes more intelligible by drawing in lines of equal intensity. The three charts represent different points along the meridian: 80° N, zenith, 80° S. In each case it can be seen that the morning hours are less bright than the evening hours. We also see that the south is always, on the average, brighter than the north, and in the south there is a clear indication that the months of September and October show the greatest degree of activity during the night.

A striking improvement in information value is obtained by taking the ratios of the intensities at 80° S with respect to 80° N. The result of this operation is shown in figure 34 where lines of equal ratio have been drawn in. The equinoctial months stand out as being those with the largest gradients along the meridian. During September at 0200 hrs., the south is brighter than the north by a factor of more than 6, on the average. In April the south

reaches peaks relative to the north at 2300 hrs., as well as 0200 hrs. The only times when the north is brighter than the south are in the early evening hours from April to September and the late morning hours from February to July.

VII. The Interpretation of the 6300A Tropical Arcs

Any attempt to interpret the many diverse facts associated with the tropical 6300A arcs should grapple with the observations on two levels. We consider that there are two categories: (a) the prime facts, and (b) the secondary features. Ultimately a complete physical picture must be concerned with both, but it is well to search for a broad explanation with respect to the prime facts and subsequently modify it in detail as the secondary observational features are brought into the discussion.

In our opinion the prime facts are:

- (a) There is a definitely established interrelationship between the 6300A tropical arcs and the F-region ionosphere,
- (b) Both the 6300A arcs and the F-region ionosphere display a region of maximum intensity or activity some 15° (\pm) from the geomagnetic equator.

The ionospherists have for some time been concerned with the problem of the anomalous behavior of the tropical F-region. It appears that the ionospheric and optical phenomena have a common origin, and we propose to consider them together.

The prime facts indicate the general direction of the inquiry. We must be concerned with charged particles moving under the influence or constraint of the earth's magnetic field. Furthermore, the charged particles move into the F-region where they affect the balance of the ionization and enter into a photochemical reaction leading to the emission of 6300A and to a lesser extent 5577A. Now, if charged particles are moving along the geomagnetic lines in the vicinity of 15° of geomagnetic latitude there must necessarily be a vertical component of motion. This leads to two, and only two, possibilities. Either the vertical component of motion is upward or downward. A priori either direction might be considered as

worthy of consideration. But two factors suggest that we should give greater weight to the downward hypothesis: (a) loss of ionization by collisions will be greater in the lower and denser parts of the atmosphere so that we may anticipate that the reservoir of ionization below the F-region would have difficulty in moving up the geomagnetic field lines; (b) gravity offers a force to propel the ionization downward.

The speculations have centered around the picture of ionization drifting sluggishly down from well above the equator to the F-region at 300 km at a latitude of about 15° . In figure 35 we show the general nature of the picture. For a terminal point at 15° latitude the equatorial origin is at 780 kilometers.

As early as 1946, Mitra offered the suggestion that the ionospheric tropical anomaly was caused by ionization in the 1000-kilometer region over the equator by the direct action of sunlight with subsequent drifting of the ionization down along the geomagnetic lines. In 1956, Martyn hypothesized that the origin of the ionization was in electrodynamic lift from the equatorial electrojet followed by a drift downward. Duncan [1960] calculated the ionization drift time to 20° magnetic latitude at various heights. In Table 8 we have assembled calculated drift times and supra-equatorial heights for pertinent latitudes based on Duncan's equations.

It is seen that a drift time of about 10 hours is required for particles originating at the equator at a height of about 780 km and arriving at the 300 km level at latitude 15° . Such a time is not inconsistent with a picture of the formation of particles at the equator, by either Mitra's or Martyn's mechanism, during the daytime and drifting to 15° latitude during the night resulting in the airglow emission.

It remains to demonstrate that such a mechanism could possibly account for the anomalous maximum in electron density located at about $\pm 15^\circ$ geomagnetic latitude. Let us assume

- 1) an electron distribution over the geomagnetic equator similar to that found by Bowles [1962];
- 2) that electrons drift down along the magnetic field lines with drift times to the 300 km level as given in Table 8;
- 3) that the number of electrons per unit time reaching the 300 km level at a given latitude depends on the number density at the equator at a height corresponding to the field line passing through the 300 km level, and to the time of drift. It is assumed that electrons are lost during the drifting period by attachment at the rate of β per second. Hence, the rate at which electrons reach the 300 km level will be proportional to $n_e e^{-\beta t}$, where n_e is the appropriate supra-equatorial electron density and t is the drift times.

Shown in figure 36 curve (a) is a typical supra-equatorial electron density profile plotted versus the latitude corresponding to the intersection with the 300 km level of the magnetic field line passing through the appropriate supra-equatorial height. Curve (b) is the function $e^{-\beta t}$ where t is the drift time from Table 8 and β has been taken at 10^{-4} , as found from trial and error to give the best result. The curve representing the product of (a) and (b) shows a pronounced maximum around 15° latitude with this choice of β .

It is of interest to note that Ratcliffe, et al. [1956] found for the loss coefficient in the F-region a value of about 10^{-4} sec^{-1} .

It should be emphasized that these ideas are of a very tentative nature and are presented primarily to stimulate discussion on the problem.

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TABLE 1

Phenomenon	Q(6300) Rayleighs	Q(5577) Rayleighs	$\frac{Q(6300)}{Q(5577)}$	References
Aurora I.B.C. 111	50,000	100,000	1/2	Chamberlain (1961)
M-arc (typical)	6,000	?	(80)*	Roach and Roach (1963)
T-arc	500	100	5	Roach, Steiger, Brown (1964)
Mid-latitude Airglow	50	250	1/5	YAO (1962)

* On only one occasion (Barbier, 1960) has there been reported a concomitant enhancement of 6300 and 5577 in an M-arc with a ratio of 80.

Table 2. Log of 6300 A Zenith Observations

<u>Night Observed</u>	<u>No. of Hours</u>	<u>Night Observed</u>	<u>No. of Hours</u>
1961			
July 18/19	8	Oct 1/2	5
19/20	8 1/2	4/5	8
20/21	7	5/6	4
		6/7	7
Aug 2/3	4	7/8	5
3/4	5	8/9	5 1/2
5/6	6 1/2	10/11	7
6/7	6 1/2	11/12	8 1/2
7/8	8 1/2		
9/10	6	Nov 3/4	9
10/11	8 1/2	4/5	9 1/2
11/12	8	5/6	5
12/13	9	6/7	9 1/2
13/14	9	7/8	10
14/15	8	8/9	10
15/16	7	9/10	10
16/17	5 1/2	10/11	10 1/2
8/31 - 9/1	4	11/12	9
		12/13	7 1/2
Sept 1/2	9 1/2	16/17	2
2/3	5 1/2	28/29	8 1/2
3/4	4	29/30	6
4/5	6 1/2	11/30 - 12/1	2
5/6	8		
6/7	6 1/2	Dec 9/10	11
7/8	8	10/11	9
10/11	4 1/2	14/15	2
11/12	8	15/16	3 1/2
12/13	9	30/31	5
13/14	8	12/31 - 1/1	6
14/15	9		
15/16	7		
16/17	6		
17/18	5 1/2		

Table 2, cont'd.

<u>Night Observed</u>	<u>No. of Hours</u>	<u>Night Observed</u>	<u>No. of Hours</u>
1962			
Jan 1/2	8	Apr 1/2	9
2/3	9 1/2	2/3	9
3/4	5	3/4	9
4/5	10	4/5	9
5/6	1 1/2	5/6	8
10/11	7 1/2	6/7	5
11/12	6	7/8	8
13/14	4	8/9	5 1/2
26/27	2 1/2	9/10	5
29/30	7	10/11	2
30/31	4	26/27	1
		28/29	2
Feb 2/3	8	29/30	7
3/4	10	4/30 - 5/1	3
4/5	10		
5/6	10 1/2	May 1/2	1/2
6/7	9 1/2	4/5	2
7/8	3	5/6	8 1/2
8/9	7	6/7	1
9/10	6 1/2	10/11	3
10/11	6	11/12	3
11/12	5 1/2	23/24	2
12/13	3	24/25	4
		25/26	5
Mar 1/2	8	26/27	5
2/3	9	28/29	7
3/4	1	29/30	8
5/6	4	5/31 - 6/1	2 1/2
6/7	8		
7/8	6 1/2	June 1/2	8
10/11	4	2/3	8 1/2
25/26	3	3/4	3 1/2
26/27	1 1/2	4/5	7
30/31	7	5/6	6
3/31 - 4/1	8	6/7	6

Table 2, cont'd.

<u>Night Observed</u>	<u>No. of Hours</u>
7/8	5
8/9	4 1/2
10/11	3
11/12	2 1/2
12/13	2
13/14	1 1/2
14/15	1 1/2
23/24	4
24/25	4 1/2
25/26	5 1/2
26/27	6
27/28	7
28/29	7 1/2
29/30	8
6/30 - 7/1	8
July 5/6	6 1/2
7/8	5
9/10	1 1/2
10/11	3 1/2
12/13	2
13/14	1
14/15	1 1/2

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

YEAR 1961 MONTH SEP IN UNITS OF RAYLEIGHS

DAY	1800	1815	1830	1845	1900	1915	1930	1945	2000	2015	2030	2045	2100	2115	2130	2145
1	0	0	0	0	0	0	0	0	176	167	192	126	0	178	0	-0
2	0	0	0	0	0	0	0	0	198	215	110	82	84	89	45	54
3	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
4	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
5	0	0	0	0	0	0	0	311	249	338	182	225	311	140	196	90
6	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
7	0	0	0	0	0	0	0	0	0	0	0	0	0	290	170	218
8	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
9	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
10	0	0	0	0	0	0	0	0	261	220	306	279	372	99	117	55
11	0	0	0	0	0	0	0	0	125	95	84	139	153	124	97	87
12	0	0	0	0	0	0	0	0	0	0	150	174	134	138	71	112
13	0	0	0	0	0	0	0	0	0	0	0	0	0	111	52	91
14	0	0	0	0	0	0	0	0	0	0	57	94	187	0	90	100
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
24	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
25	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
26	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
29	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
30	0	0	0	0	0	0	0	0	0	0	0	0	0	77	155	182

5300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

IN UNITS OF RAYLEIGHS

YEAR 1961 MONTH NOV

DAY 1800 1815 1830 1845 1900 1915 1930 1945 2000 2015 2030 2045 2100 2115 2130 2145

1	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
2	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
3	0	0	0	172	279	166	199	103	157	130	115	128	132	135	185
4	0	0	0	0	0	0	0	182	180	171	137	120	62	64	72
5	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
6	0	0	0	0	0	0	0	134	156	138	106	114	84	123	110
7	0	0	0	0	276	182	117	150	101	122	190	150	180	100	177
8	0	0	0	194	359	160	190	207	240	128	135	111	50	144	109
9	0	0	0	374	245	169	16	139	85	101	136	112	106	115	122
10	0	0	0	0	0	0	0	174	37	107	113	103	94	5	120
11	0	0	0	0	0	0	0	0	0	0	0	0	199	103	150
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
24	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
25	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
26	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	0	0	0	0	207	201	151	133	89	54	73	71	44	64	63
29	0	0	0	191	190	174	123	86	70	85	105	72	61	79	86
30	0	0	0	0	147	218	133	109	82	87	113	104	50	97	

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

YEAR 1961 MONTH NOV IN UNITS OF RAYLEIGHS

DAY 2200 2215 2230 2245 2300 2315 2330 2345 0000 0015 0030 0045 0100 0115 0130 0145

1	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
2	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
3	83	0	79	0	126	118	148	137	109	88	85	84	57	81	74
4	51	90	59	51	61	122	18	25	18	163	113	122	111	63	133
5	0	120	182	187	114	117	46	103	120	63	89	136	77	45	0
6	111	206	163	149	179	113	130	118	134	151	181	177	172	128	108
7	117	115	100	82	57	141	74	55	66	90	100	123	107	73	35
8	125	122	81	96	123	114	101	114	117	127	110	130	95	83	67
9	137	101	107	82	106	120	75	61	79	45	31	86	80	64	31
10	71	95	87	85	69	87	60	71	59	82	78	61	54	59	65
11	96	74	64	86	61	45	90	95	72	106	160	220	109	134	132
12	53	74	78	154	56	98	62	70	51	46	29	39	54	59	42
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
24	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
25	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
26	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	64	85	65	69	134	86	50	45	0	0	0	0	0	0	0
29	70	60	37	35	32	58	22	51	57	56	38	0	0	0	0
30	100	107	159	95	139	112	110	152	172	145	135	115	106	113	115

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

IN UNITS OF RAYLEIGHS

YEAR 1961 MONTH DEC

DAY 2200 2215 2230 2245 2300 2315 2330 2345 0000 0015 0030 0045 0100 0115 0130 0145

1	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
2	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
3	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
4	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
5	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
6	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
7	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
8	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
9	55	64	54	48	48	54	46	31	26	39	26	15	40	17	28
10	108	110	87	67	94	59	107	74	51	51	52	90	39	44	53
11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
24	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
25	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
26	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
29	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
30	61	78	76	76	80	54	38	80	45	59	55	41	38	26	20
31	70	76	97	62	55	73	45	46	57	43	0	0	0	0	0

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

IN UNITS OF RAYLEIGHS

YEAR 1962 MONTH MAR

DAY	1800	1815	1830	1845	1900	1915	1930	1945	2000	2015	2030	2045	2100	2115	2130	2145
1	0	0	0	0	0	0	0	110	272	88	172	118	84	79	71	78
2	0	0	0	0	0	0	0	272	140	237	125	97	63	71	86	79
3	0	0	0	0	0	0	0	222	154	132	80	103	0	0	0	0
4	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
5	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
6	0	0	0	0	0	0	0	0	0	72	74	97	49	50	47	52
7	0	0	0	0	0	0	0	0	0	0	0	6456	6151	5405	6128	0
8	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
9	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
10	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
24	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
25	0	0	0	0	0	0	0	0	424	283	521	252	257	150	256	0
26	0	0	0	0	0	0	0	0	785	0	0	0	0	0	0	0
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
29	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
30	0	0	0	0	0	0	0	0	265	283	312	244	221	200	91	158
31	0	0	0	0	0	0	0	0	187	134	78	84	85	60	70	25

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

YEAR 1962 MONTH MAR IN UNITS OF RAYLEIGHS

DAY 2200 2215 2230 2245 2300 2315 2330 2345 0000 0015 0030 0045 0100 0115 0130 0145

1	51	47	44	36	54	36	44	18	5	19	38	56	35	43	37	45
2	38	43	63	49	68	55	47	41	35	25	48	45	46	44	44	33
3	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
4	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
5	0	0	0	0	84	56	46	0	0	0	0	0	0	0	0	0
6	64	56	57	0	0	0	51	96	36	47	34	43	58	57	40	37
7	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
8	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
9	0	0	0	0	0	0	0	0	0	46	48	50	60	49	27	39
10	0	0	0	0	0	88	74	55	50	0	0	0	0	0	0	0
11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
24	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
25	194	152	76	80	0	0	0	0	0	0	0	0	0	0	0	0
26	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
29	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
30	35	134	24	58	5	51	19	65	67	73	87	84	89	62	80	72
31	41	74	55	67	52	58	54	62	64	74	98	89	74	78	83	72

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

IN UNITS OF RAYLEIGHS

YEAR 1962 MONTH MAR

DAY 0200 0215 0230 0245 0300 0315 0330 0345 0400 0415 0430 0445 0500 0515 0530 0545

1	43	42	36	59	32	27	27	0	0	0	0	0	0	0	0
2	35	37	38	40	48	42	11	13	0	28	25	17	0	0	0
3	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
4	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
5	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
6	28	41	31	37	56	38	0	0	42	38	44	34	29	40	33
7	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
8	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
9	42	34	42	27	46	0	0	0	0	0	0	0	0	0	0
10	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
24	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
25	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
26	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
29	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
30	0	128	88	80	48	151	0	0	0	0	0	0	0	0	0
31	114	57	37	36	57	32	28	29	0	0	0	0	0	0	0

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

YEAR 1962 MONTH APR

IN UNITS OF RAYLEIGHS

DAY 2200 2215 2230 2245 2300 2315 2330 2345 0000 0015 0030 0045 0100 0115 0130 0145

1	21	39	19	14	19	16	18	25	33	26	23	56	30	47	22	36
2	33	53	51	33	40	28	41	33	38	42	60	98	76	54	81	25
3	64	29	15	18	47	41	44	31	37	49	49	23	13	24	9	14
4	42	36	16	25	16	29	23	36	31	25	33	42	40	33	38	0
5	47	43	20	27	13	39	42	47	47	56	63	43	25	32	29	86
6	0	0	0	0	0	0	0	0	0	38	44	60	34	26	28	29
7	0	0	39	37	26	16	32	14	22	32	30	32	26	39	0	34
8	0	0	0	0	0	0	0	19	13	15	26	26	22	62	0	52
9	0	0	0	0	0	0	0	0	0	43	41	51	26	0	44	17
10	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
24	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
25	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
26	35	96	57	29	114	76	113	151	141	0	235	270	317	0	0	0
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	0	0	0	0	0	0	0	0	0	0	0	50	67	80	56	30
29	48	38	40	37	0	34	28	23	61	83	36	24	21	56	46	49
30	0	0	0	0	0	0	0	0	0	0	0	0	0	92	64	61

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

IN UNITS OF RAYLEIGHS

YEAR 1962 MONTH APR

DAY 0200 0215 0230 0245 0300 0315 0330 0345 0400 0415 0430 0445 0500 0515 0530 0545

1	0	45	39	38	44	56	54	39	25	54	43	40	32	0	0	0
2	142	56	57	73	48	31	41	33	28	23	20	26	22	0	0	0
3	7	20	13	26	50	57	52	39	28	47	54	29	21	0	0	0
4	31	47	44	33	45	46	34	36	20	16	17	23	47	0	0	0
5	39	37	32	41	21	24	47	47	71	112	84	79	72	0	0	0
6	33	47	35	55	68	54	33	42	46	36	41	40	47	0	0	0
7	22	47	56	62	62	57	35	44	50	35	34	35	40	0	0	0
8	48	40	75	51	63	39	36	31	29	29	26	33	54	0	0	0
9	20	24	21	29	31	22	16	37	39	48	52	46	64	0	0	0
10	0	0	0	34	29	30	36	52	27	36	55	40	55	0	0	0
11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
24	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
25	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
26	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	90	59	70	74	106	0	0	0	0	0	0	0	0	0	0	0
29	43	40	47	55	32	36	20	0	0	0	0	0	0	0	0	0
30	74	117	72	35	46	73	80	51	79	97	0	0	0	0	0	0

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

IN UNITS OF RAYLEIGHS

YEAR 1962 MONTH MAY

DAY 1800 1815 1830 1845 1900 1915 1930 1945 2000 2015 2030 2045 2100 2115 2130 2145

1	0	0	0	0	0	0	0	200	259	0	0	0	0	0	0
2	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
3	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
4	0	0	0	0	0	0	0	315	273	0	0	248	81	63	
5	0	0	0	0	0	0	0	0	97	67	72	53	48		
6	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
7	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
8	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
9	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
10	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	0	29	35	47		
24	0	0	0	0	0	0	0	325	264	181	207	113	106	37	
25	0	0	0	0	0	0	0	141	351	149	140	80	97	107	
26	0	0	0	0	0	0	0	200	137	167	228	72	80		
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	0	0	0	0	0	0	0	0	339	29	40	850	0		
29	0	0	0	0	0	0	0	191	335	98	66	99	93	85	
30	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
31	0	0	0	0	0	0	0	0	0	0	36	37	97		

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

IN UNITS OF RAYLEIGHS

YEAR 1962 MONTH MAY

DAY 2200 2215 2230 2245 2300 2315 2330 2345 0000 0015 0030 0045 0100 0115 0130 0145

1	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
2	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
3	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
4	66	0	0	0	0	0	0	34	32	35	50	37	47	56	39
5	21	38	96	57	33	72	109	0	83	36	0	50	0	40	66
6	0	0	0	0	0	0	0	0	0	29	42	37	0	0	0
7	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
8	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
9	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
10	0	0	0	0	0	0	0	0	0	0	0	0	0	67	67
11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	105
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	78	46	0	47	27	0	0	0	0	0	0	0	0	0	0
24	66	43	67	67	32	61	24	23	19	0	0	0	0	0	0
25	54	0	41	17	22	27	38	73	63	82	79	58	98	0	0
26	59	0	29	36	38	27	53	79	47	47	37	73	68	80	66
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	142	212	149	0	144	0	37	0	17	13	151	0	0	31	0
29	10	64	66	62	117	80	108	97	88	65	75	161	60	78	92
30	0	0	0	0	30	255	17	0	26	0	0	8	165	39	60
31	0	54	53	83	11	48	50	75	0	0	0	0	0	0	0

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

IN UNITS OF RAYLEIGHS

YEAR 1962 MONTH JUN

DAY 1800 1815 1830 1845 1900 1915 1930 1945 2000 2015 2030 2045 2100 2115 2130 2145

1	0	0	0	0	0	0	0	0	0	239	0	210	41	32	16
2	0	0	0	0	0	0	0	0	0	197	533	345	194	32	
3	0	0	0	0	0	0	0	0	0	331	199	66	52	112	
4	0	0	0	0	0	0	0	0	0	0	0	188	6	0	0
5	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
6	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
7	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
8	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
9	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
10	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	51	25	8	0	9	
24	0	0	0	0	0	0	0	0	0	172	144	117	115	91	
25	0	0	0	0	0	0	0	0	0	447	437	211	337	323	207
26	0	0	0	0	0	0	0	0	0	128	101	167	61	93	58
27	0	0	0	0	0	0	0	0	0	45	316	164	0	229	
28	0	0	0	0	0	0	0	0	0	71	80	76	-0	57	106
29	0	0	0	0	0	0	0	0	0	65	56	41	19	34	
30	0	0	0	0	0	0	0	0	0	5	52	0	7	0	

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

IN UNITS OF RAYLEIGHS

YEAR 1962 MONTH JUN

DAY 2200 2215 2230 2245 2300 2315 2330 2345 0000 0015 0030 0045 0100 0115 0130 0145

1	5393	19	13	44	16	58	0	110	162	0	41	0	8	87	0	46
2	218	81	19	53	23	20	21	86	4	8	37	30	87	26	0	19
3	23	37	23	20	12	16	18	0	0	0	0	0	0	0	0	0
4	46	73	7	218	72	8	55	71	33	46	0	8	11	29	16	17
5	0	43	24	37	0	0	53	0	-0	0	24	0	28	9	16	43
6	0	0	0	0	5	95	20	25	2	330	0	0	0	19	11	26
7	0	0	0	0	0	0	0	63	85	61	74	0	9	59	75	71
8	0	0	0	0	0	0	0	0	0	0	37	0	65	46	10	0
9	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
10	0	0	0	0	0	0	0	0	0	0	0	0	0	0	73	98
11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	7	10	0	67	0	0	0	0	0	0	0	0	0	0	0	0
24	49	124	86	58	95	22	64	93	0	50	95	119	65	0	0	0
25	167	205	275	146	107	149	69	37	104	62	56	139	132	68	101	84
26	39	59	43	68	75	29	21	53	243	37	30	59	39	74	77	69
27	84	21	30	21	34	11	75	35	14	66	24	118	130	72	103	100
28	82	72	27	0	39	30	33	174	189	122	38	41	64	83	70	84
29	0	26	50	37	31	50	43	37	58	10	51	29	19	21	34	0
30	62	0	0	0	0	0	3	0	0	0	0	0	40	0	0	7

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

YEAR 1962 MONTH JUN IN UNITS OF RAYLEIGHS

DAY 0200 0215 0230 0245 0300 0315 0330 0345 0400 0415 0430 0445 0500 0515 0530 0545

1	0	125	18	0	129	0	21	29	102	0	27	0	0	0	0
2	7	0	106	4	99	53	19	7	13	100	-0	0	0	0	0
3	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
4	69	0	95	40	15	199	90	73	68	31	65	0	0	0	0
5	22	0	0	40	8	56	119	10	14	58	0	0	0	0	0
6	4	22	26	0	0	0	0	0	13	5	0	0	0	0	0
7	14	16	58	49	25	81	39	52	145	166	162	0	0	0	0
8	14	89	40	3	56	0	90	75	118	76	0	0	0	0	0
9	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
10	68	92	123	23	170	205	66	115	61	77	46	0	0	0	0
11	0	41	121	19	69	3	13	52	14	29	45	0	0	0	0
12	0	0	0	0	129	10	5	44	0	17	16	0	0	0	0
13	0	0	0	0	0	13	48	40	41	57	54	0	0	0	0
14	0	0	0	0	0	0	0	0	228	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
24	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
25	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
26	56	71	44	0	0	0	0	0	0	0	0	0	0	0	0
27	49	108	75	191	0	-0	0	0	0	0	0	0	0	0	0
28	79	119	170	184	121	84	50	98	19	0	0	0	0	0	0
29	3	14	76	47	86	37	48	56	76	58	102	0	0	0	0
30	60	143	46	22	69	179	32	36	0	29	27	0	0	0	0

6300A AIRGLOW ZENITH INTENSITIES
HALEAKALA, HAWAII

IN UNITS OF RAYLEIGHS

YEAR 1962 MONTH JUL

DAY 1800 1815 1830 1845 1900 1915 1930 1945 2000 2015 2030 2045 2100 2115 2130 2145

1	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
2	0	0	0	0	0	0	0	0	0	0	147	104	75	0	0
3	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
4	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
5	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
6	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
7	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
8	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
9	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
10	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
11	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
12	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
13	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
14	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
15	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
16	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
17	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	0	0	0	0	0	0	0	0	44	0	0	0	0	0	0
19	0	0	0	0	0	0	0	0	0	0	0	231	236	443	0
20	0	0	0	0	0	0	0	0	0	0	0	0	229	312	111
21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
22	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
23	0	0	0	0	0	0	0	0	0	94	79	48	177	315	0
24	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
25	0	0	0	0	0	0	0	0	0	0	0	0	0	0	9
26	0	0	0	0	0	0	0	0	0	116	78	42	355	44	0
27	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
28	0	0	0	0	0	0	0	0	175	92	52	76	112	43	0
29	0	0	0	0	0	0	0	0	0	140	59	38	33	58	43
30	0	0	0	0	0	0	0	0	0	0	41	47	33	67	0
31	0	0	0	0	0	0	0	0	163	218	0	0	0	0	0

Table 5: Observations Observed and Calculated Q(6300) Together With

Selected Ionospheric and Geomagnetic Data for 93 Nights

DATE		MEAN* IONOSPHERIC PARAMETERS				MEAN* Q(6300) (rayleighs)		Kp
Month	Day	f _o F ₂ (Mc/s)	h'F ₂ (km)	E	Obs.	Calc.		
1961								
9	5/6	5.1	260	5.88	65	77	1-, 1+	
9	11/12	9.3	228	50.26	259	248	5 _o , 2+	
10	5/6	4.5	240	11.8	73	157	1+, 2-, 1+	
10	8/9	3.2	244	3.97	89	85	1 _o , 1-, 0+	
10	11/12	5.9	236	19.62	94	154	4 _o , 3-, 3-	
10	19/20	3.5	241	5.63	85	80	1+, 2-, 0+	
10	24/25	3.5	233	5.52	100	101	1 _o , 1 _o , 2+	
11	4/5	3.8	245	6.46	94	110	3 _o , 3 _o , 5-	
11	5/6	3.9	238	9.08	105	108	5-, 4 _o , 2+	
11	6/7	5.0	225	13.48	150	155	4+, 5 _o , 4+	
11	7/8	3.7	237	6.99	72	87	4 _o , 2 _o , 3+	
11	8/9	4.2	226	9.40	89	96	4-, 2+, 2+	
11	9/10	2.8	236	5.13	59	70	1 _o , 1 _o , 1 _o	
11	10/11	2.7	234	3.58	63	67	1-, 0+, 0+	
11	11/12	4.8	232	10.52	103	113	3+, 2+, 3 _o	
11	12/13	3.0	237	4.20	54	58	1-, 1-, 0 _o	
11	16/17	2.6	209	5.60	79	84	1 _o , 1-, 4 _o	
11	19/20	2.8	232	4.28	100	112	4-, 3-, 4-	
11	27/28	3.2	250	3.36	63	76	0 _o , 0+, 0+	
11	28/29	3.1	254	2.85	38	41	1 _o , 0+, 0+	
11	29/30	3.3	239	4.36	44	45	1 _o , 0 _o , 0+	
12	9/10	2.5	248	2.84	39	47	0+, 2 _o , 2-	
12	10/11	3.3	257	5.00	60	93	4-, 3 _o , 1+	
12	15/16	2.8	230	3.88	34	48	1+, 1 _o , 1-	
12	30/31	2.6	285	1.37	41	41	1+, 1+, 1+	
12	31/1	2.2	304	1.25	36	30	1 _o , 1-, 1-	

* MEANS are for 5 hrs., 2200-0300

Table 5, cont'd.

DATE		MEAN IONOSPHERIC PARAMETERS			MEAN Q(6300) (rayleighs)	Kp
Month	Day	f _o F ₂ (Mc/s)	h'F ₂ (km)	E		
1962						
1	1/2	2.5	287	1.65	29	30, 3+, 2-
1	2/3	2.3	270	1.62	32	00, 1-, 1-
1	3/4	2.0	250	0.88	15	0+, 1-, 00
1	29/30	2.9	248	2.33	35	2+, 2-, 20
1	30/31	2.9	236	3.71	50	0+, 00, 00
2	2/3	3.2	234	4.87	49	1+, 0+, 1-
2	3/4	3.7	223	7.95	63	2-, 3+, 50
2	4/5	3.4	250	4.59	45	0+, 1-, 2-
2	5/6	2.9	228	5.48	48	1-, 10, 1-
2	6/7	3.1	242	3.88	42	4-, 40, 3+
2	7/8	2.8	236	3.43	34	1-, 2-, 2-
2	8/9	2.6	257	3.26	31	10, 10, 20
2	9/10	2.9	243	3.05	33	0+, 0+, 10
2	10/11	2.5	262	1.66	28	0+, 3-, 3-
2	11/12	3.2	257	2.79	28	30, 3-, 20
2	12/13	2.3	258	1.56	36	2-, 3+, 3-
3	1/2	3.6	234	6.12	31	20, 1-, 3-
3	2/3	3.6	222	7.44	52	1-, 10, 3+
3	5/6	3.4	246	4.79	67	4+, 4+, 6-
3	6/7	3.4	229	5.31	45	30, 2-, 2+
3	7/8	3.8	237	6.02	58	10, 00, 0+
3	10/11	3.2	244	4.18	61	2+, 30, 2-
3	31/1	4.9	233	11.24	64	4-, 4-, 2+
4	1/2	3.5	251	4.04	35	0+, 10, 10
4	2/3	4.3	247	8.18	51	4-, 20, 2-
4	3/4	4.9	252	7.12	35	4-, 40, 2-
4	4/5	3.5	252	3.93	33	1-, 2-, 10
4	7/8	3.9	276	3.97	36	30, 3-, 4-
4	8/9	4.6	264	5.82	31	30, 2-, 30
4	9/10	4.4	290	2.43	20	20, 4+, 50
4	26/27	6.5	241	24.65	98	2-, 2-, 1+

Table 5, cont'd.

DATE Month Day	MEAN IONOSPHERIC PARAMETERS			MEAN Q(6300) (rayleighs) Obs. Calc.	Kp	
	f _o F ₂ (Mc/s)	h'F ₂ (km)	E			
1962						
4 28/29	5.3	261	6.21	73	78	1 ₀ , 1 ₊ , 2 ₊
4 29/30	5.6	282	4.75	35	100	1 ₋ , 1 ₋ , 0 ₊
4 30/1	5.3	260	8.27	73	76	1 ₀ , 1 ₀ , 1 ₋
5 5/6	6.9	267	9.43	69	69	1 ₊ , 3 ₊ , 4 ₊
5 10/11	6.9	267	9.30	76	68	1 ₊ , 0 ₊ , 3 ₀
5 11/12	7.2	269	9.34	84	85	1 ₊ , 1 ₀ , 0 ₊
5 23/24	6.8	261	10.90	93	101	0 ₊ , 0 ₊ , 0 ₀
5 24/25	6.1	260	8.32	48	51	0 ₀ , 0 ₊ , 0 ₊
5 25/26	7.7	257	15.85	73	84	1 ₋ , 1 ₋ , 1 ₀
5 26/27	7.8	288	6.87	62	58	3 ₋ , 2 ₊ , 3 ₀
5 28/29	6.3	265	12.84	22	27	2 ₊ , 2 ₋ , 1 ₊
5 29/30	6.9	245	17.80	86	94	1 ₀ , 0 ₊ , 1 ₀
5 31/1	7.1	271	8.29	54	54	3 ₀ , 3 ₋ , 2 ₀
6 1/2	7.1	250	15.10	33	34	1 ₋ , 2 ₋ , 1 ₋
6 4/5	7.2	265	12.22	25	30	2 ₊ , 2 ₀ , 1 ₀
6 6/7	5.8	288	4.24	11	11	4 ₋ , 3 ₋ , 2 ₀
6 7/8	5.7	262	8.12	88	72	1 ₊ , 1 ₋ , 1 ₋
6 8/9	7.7	274	11.12	77	107	3 ₊ , 4 ₋ , 4 ₋
6 10/11	6.7	268	9.44	68	82	2 ₀ , 1 ₊ , 2 ₋
6 11/12	7.7	250	19.61	91	136	2 ₊ , 2 ₊ , 3 ₋
6 12/13	6.2	261	9.27	60	67	2 ₊ , 2 ₋ , 1 ₋
6 13/14	8.2	235	31.18	148	152	2 ₋ , 2 ₋ , 2 ₋
6 14/15	5.8	271	6.27	8	9	2 ₀ , 2 ₀ , 3 ₀
6 15/16	5.5	264	6.66	8	8	1 ₀ , 1 ₀ , 1 ₊
6 23/24	7.3	259	13.2	84	97	2 ₀ , 2 ₋ , 1 ₊
6 24/25	7.3	253	15.20	110	99	3 ₋ , 1 ₊ , 1 ₊
6 25/26	6.3	243	17.50	103	125	2 ₀ , 1 ₊ , 1 ₊
6 26/27	7.0	---	---	40	40	4 ₋ , 3 ₀ , 4 ₀
6 27/28	6.2	259	10.32	62	74	3 ₊ , 4 ₋ , 4 ₋

Table 5, cont'd.

DATE		MEAN IONOSPHERIC PARAMETERS			MEAN Q(6300)		Kp
Month	Day	f _o F ₂ (Mc/s)	h'F ₂ (km)	E	Obs.	Calc.	
1962							
6	28/29	5.2	259	6.60	89	83	3 _o , 3-, 3+
6	29/30	5.3	285	4.18	41	61	3+, 3 _o , 2 _o
6	30/1	6.1	250	11.44	84	100	3-, 3 _o , 2+
7	5/6	6.4	274	6.81	61	72	3-, 4-, 2 _o
7	10/11	5.4	247	11.51	87	91	3-, 2 _o , 3-
7	12/13	5.2	258	6.81	39	38	2-, 1+, 3-
7	13/14	6.4	258	12.10	84	104	3-, 2-, 2-

Table 6: Correlation Between Observed and Calculated
Q(6300) for Various Values of Scale Height

Yr.	Date Mo.	Night	Correlation Coefficient			No. of Observations Used
			H = 30 km	40 km	50 km	
61	9	5/6	0.979	0.980	0.978	39
		11/12	761	707	660	39
	10	5/6	943	946	948	24
		8/9	810	869	904	27
		11/12	934	969	984	32
		19/20	920	932	925	11
		24/25	911	904	899	15
	11	4/5	939	944	945	42
		5/6	954	954	950	25
		6/7	934	927	913	35
		7/8	908	941	955	36
		8/9	944	942	938	44
		9/10	888	887	879	39
		10/11	872	894	899	45
		11/12	973	970	958	44
		12/13	572	648	696	35
		16/17	578	689	743	24
		19/20	756	770	776	28
		27/28	350	340	316	38
		28/29	938	937	931	40
		29/30	972	973	972	32
		12	9/10	879	888	887
	10/11		587	608	621	19
	15/16		010	120	231	8
	30/31		773	832	861	36
	31/1		882	915	930	19
	62	1	1/2	724	772	794
2/3			941	950	953	41
3/4			905	891	877	7
29/30			962	965	968	11
30/31			943	951	956	8
2		2/3	436	521	575	30
		3/4	819	802	782	44
		4/5	814	778	744	41
		5/6	562	597	600	41
		6/7	487	545	567	41
		7/8	937	957	966	23
		8/9	515	501	492	25

Table 6, cont'd.

Yr.	Date Mo.	Night	Correlation Coefficient			No. of Observations Used
			H = 30 km	40 km	50 km	
		9/10	0.620	0.590	0.564	34
		10/11	667	694	705	34
		11/12	835	859	875	16
		12/13	939	950	951	24
3		1/2	722	847	899	37
		2/3	816	884	910	39
		5/6	919	944	959	4
		6/7	074	187	271	33
		7/8	277	243	221	7
		10/11	797	832	853	26
		31/1	255	236	221	23
4		1/2	772	778	781	39
		2/3	417	416	413	33
		3/4	829	831	831	40
		4/5	953	958	961	31
		7/8	713	740	755	34
		8/9	904	907	902	40
		9/10	725	677	608	31
		26/27	037	056	075	13
		28/29	727	622	499	3
		29/30	223	172	125	14
		30/1	-0.055	-0.098	-0.127	14
5		5/6	0.660	0.645	0.627	28
		10/11	954	940	930	9
		11/12	840	853	863	3
		23/24	683	594	474	33
		24/25	929	943	948	32
		25/26	830	849	850	31
		26/27	810	844	860	21
		28/29	920	937	941	26
		29/30	383	437	477	25
		31/1	420	419	419	7
6		1/2	1.000	1.000	1.000	2
		4/5	0.667	0.634	0.590	27
		6/7	863	865	860	33
		7/8	911	928	936	34
		8/9	710	703	688	34
		10/11	368	343	310	36
		11/12	867	849	834	5
		12/13	513	583	624	26
		13/14	943	963	971	33
		14/15	682	709	718	32
		15/16	906	932	937	34
		23/24	824	853	856	6
		24/25	923	858	789	7
		25/26	989	995	997	7
		26/27	995	997	998	3
		27/28	705	725	732	10
		28/29	837	763	669	8
		29/30	431	444	444	9
		30/1	953	943	932	10
7		5/6	095	037	-0.016	8
		10/11	938	958	0.964	36
		12/13	962	966	967	11
		13/14	715	726	728	31
Ave			0.735	0.745	0.743	

Table 7

SEASONAL VARIATION IN POSITION
OF 6300 EVENING MAXIMUM ALONG MERIDIAN

GROUP	No. of Nights	Mean Date	Dec. of Sun	6300 Evening Maximum		
				Position km	Lat.	L
1	16	1961 June 9	22.90	+ 645	26.6	1.28
2	15	July 13	21.90	+ 670	26.8	1.28
3	14	Aug. 12	15.12	+ 100	21.6	1.19
4	16	Sept. 9	5.33	- 225	18.7	1.14
5	13	Oct. 13	- 7.60	- 530	15.9	1.11
6	12	Nov. 10	-17.02	- 960	12.0	1.07
7	10	1961 Dec. 7	-22.51	- 950	12.1	1.07
8	6	1962 Jan. 2	-22.98	- 1090	10.8	1.07
9	11	Feb. 6	-15.83	- 850	13.0	1.08
10	4	Mar. 6	- 5.93	- 650	14.8	1.10
11	10	Apr. 2	+ 4.65	- 220	18.7	1.14
12	4	May 4	+15.77	- 100	19.8	1.16
13	18	June 6	+22.58	+ 325	23.7	1.22

Table 8. Electron Drift Times to 300 kilometers

Latitudes (Geomagnetic)	Equatorial Height (kilometers)	Drift Time (hours)	
		500 to 300 km	Equator to 300 km
25°	1750	4.48	5.5
24	1622	4.72	5.7
23	1502	5.00	6.0
22	1389	5.32	6.3
21	1283	5.68	6.7
20	1184	6.10	7.1
19	1091	6.53	7.5
18	1004	7.09	8.1
17	924	7.75	8.7
16	848	8.47	9.5
15	779	9.43	10.4
14	715	10.6	11.6
13	656	12.2	13.2
12	601	14.1	15.1
11	553	16.9	17.9
10	507	21.5	22.5
9	467	28.6	29.6
8	432	42.5	43.5
7	400	83.3	84.3
6	374		
5	351		

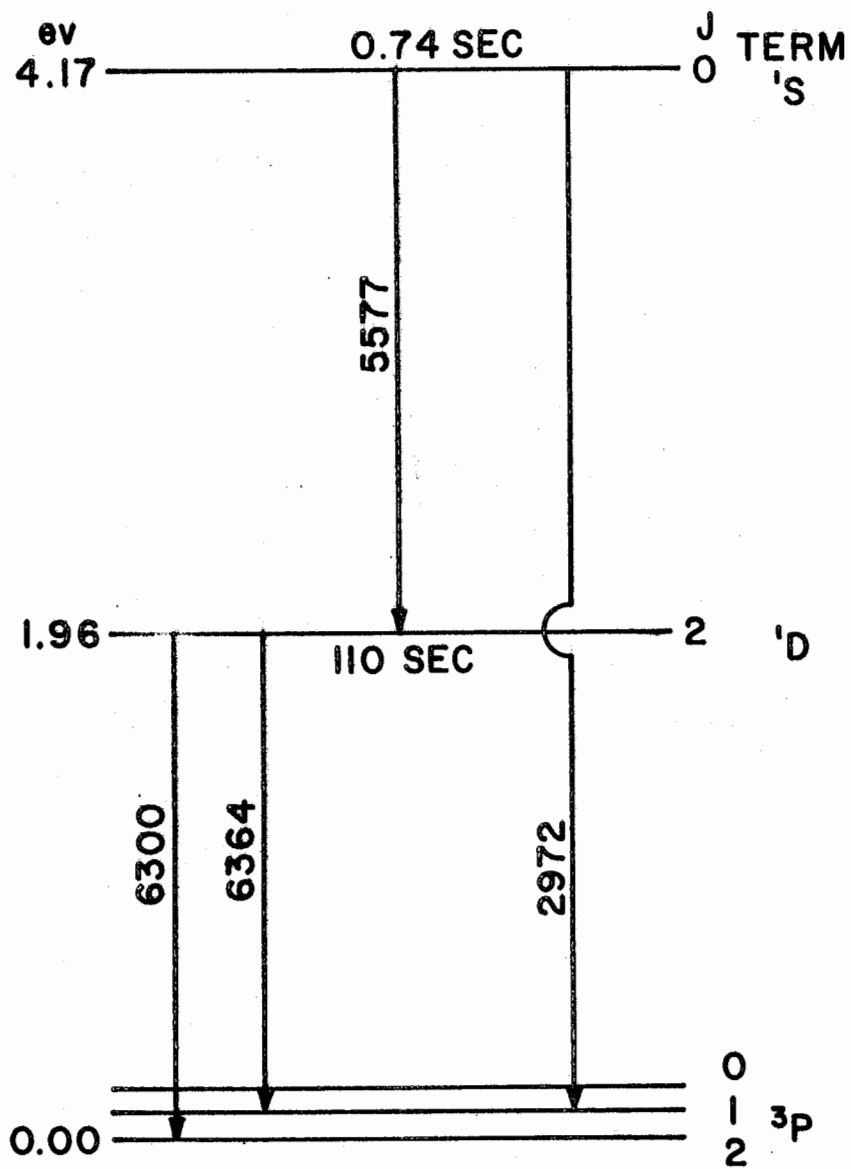


Figure 1

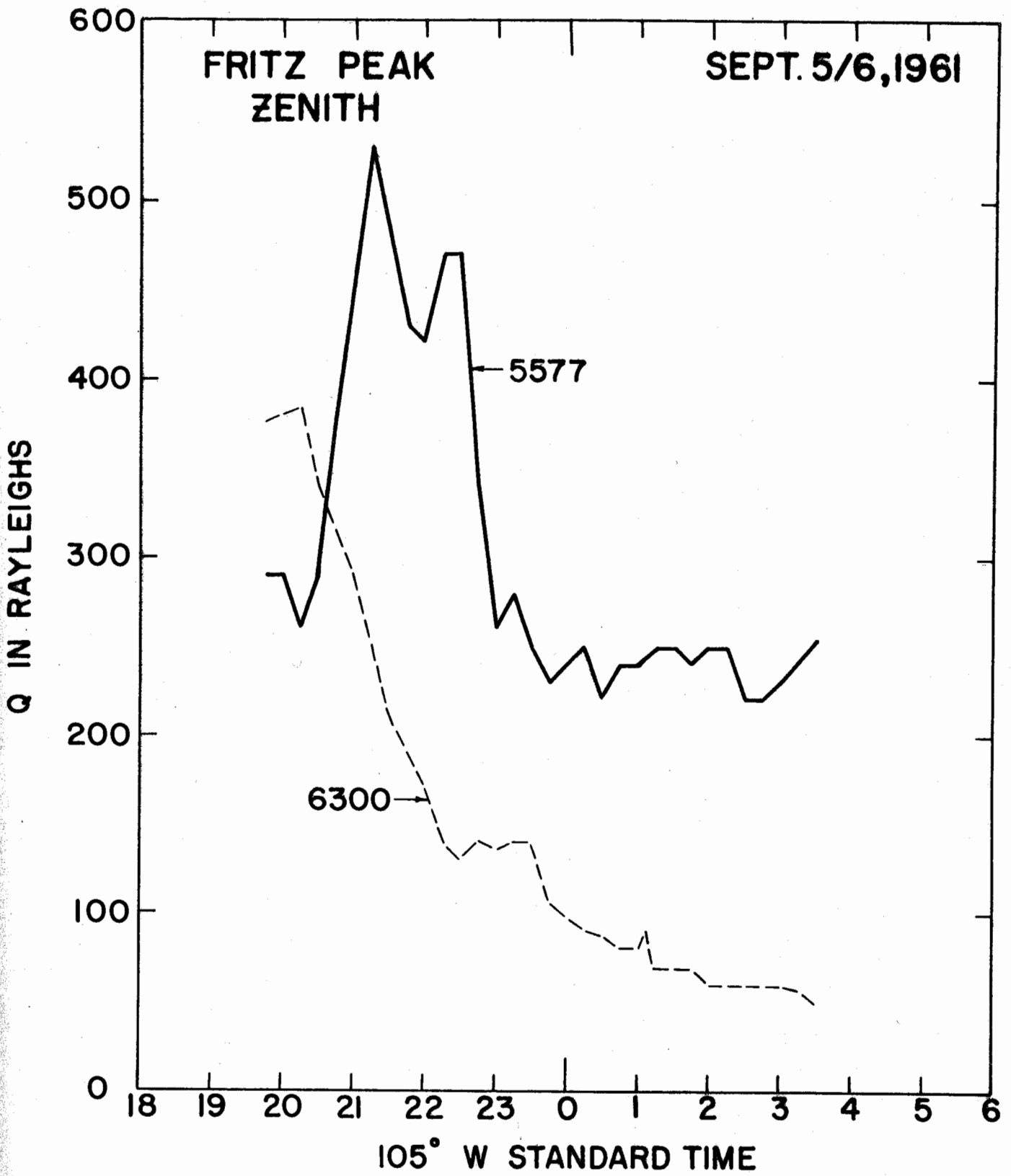


Figure 2

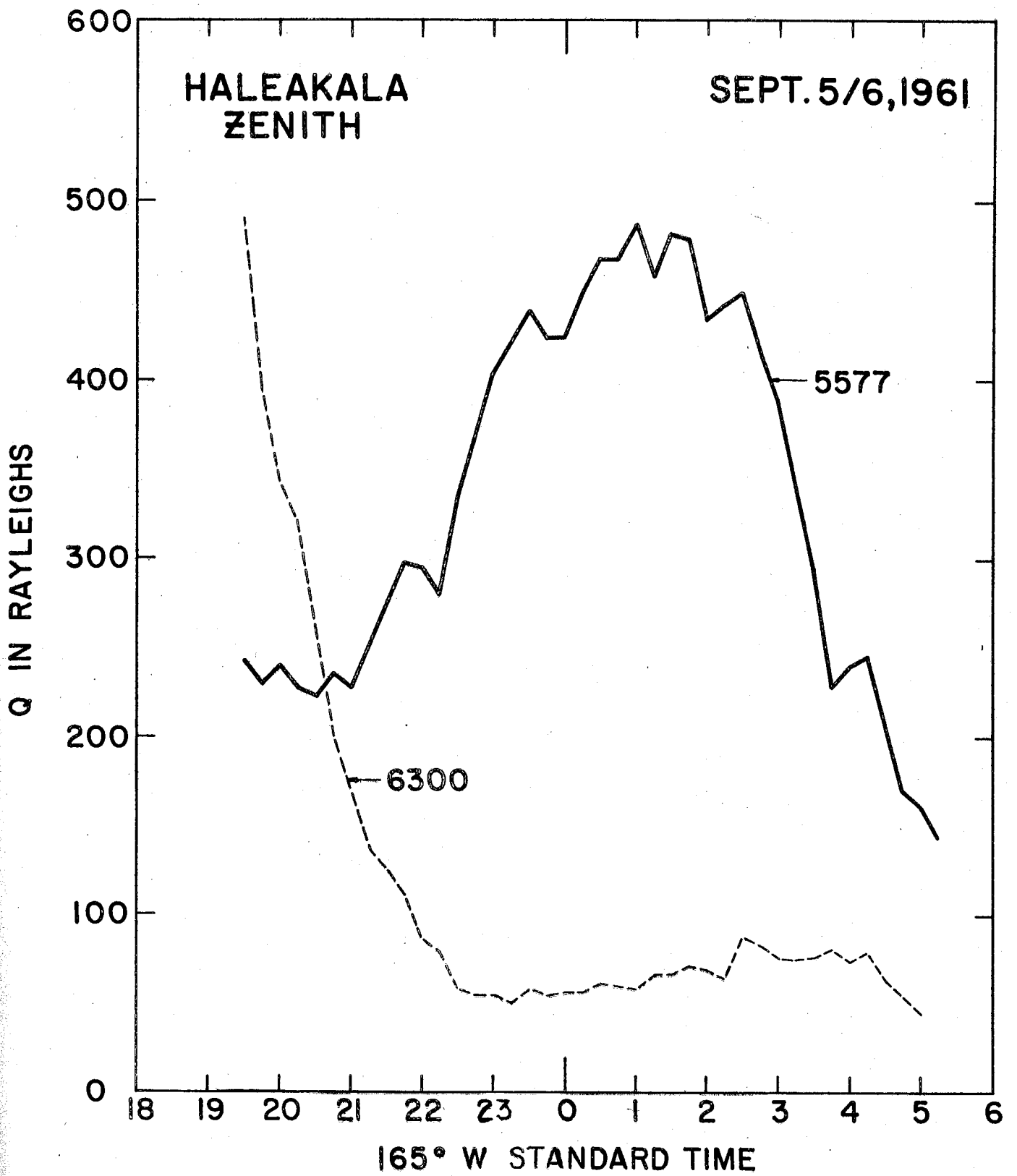


Figure 3

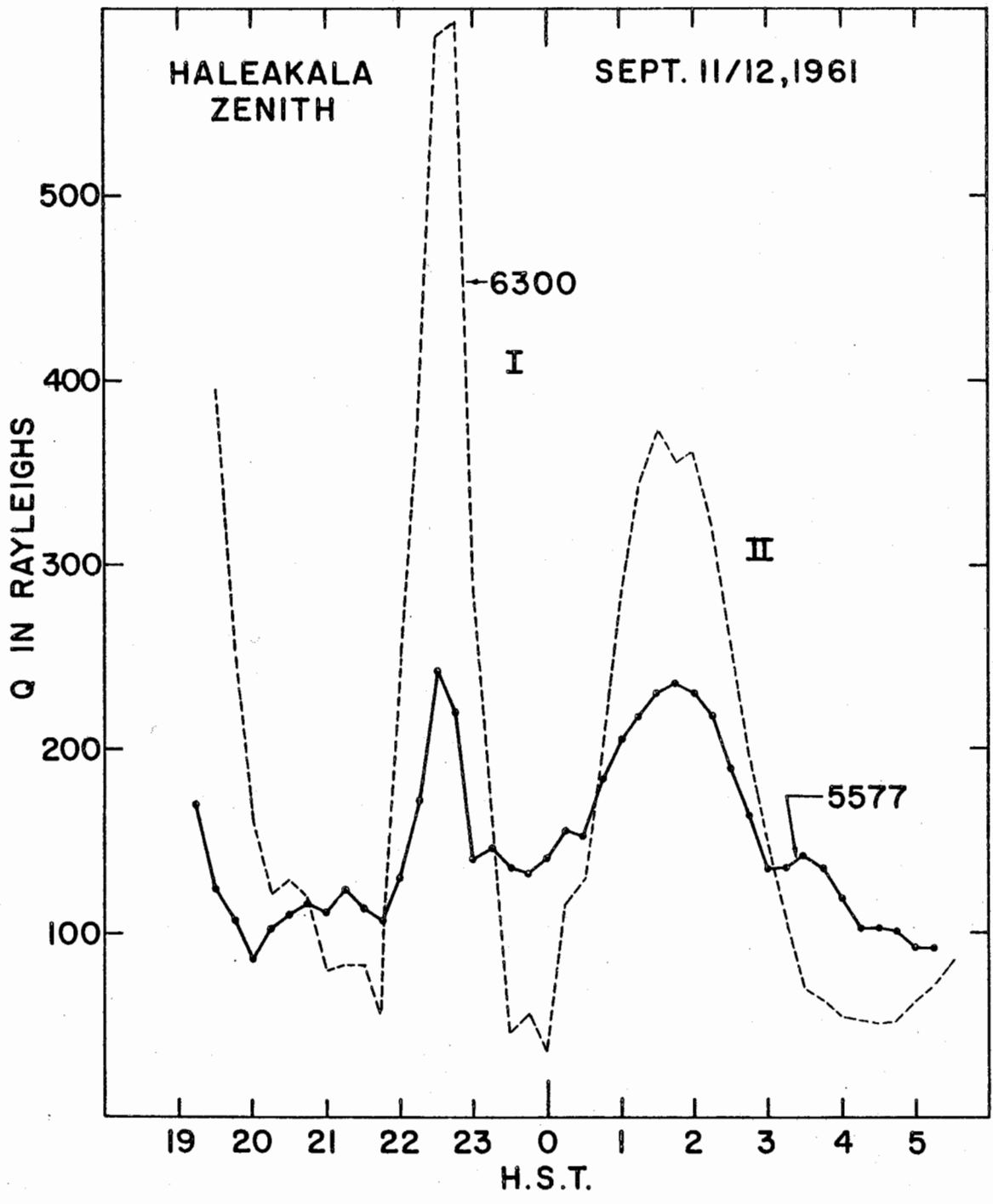


Figure 4

Table 3

PROGRAM LZIB63										STA 20															
Q= C.46										REL SENS=200000.															
SEN	J	TYPE	ALT	C	YR	MO	DA	MST	N	E	S	W	Z												
254	333	330	80	3	61	9	11	2225	79	60	137	165	139	289	290	224	157	140	140	394	542	286	214	175	585
			75						108	67	228	260	356	512	386	275	226	121	397	525	467	330	295	111	
			70						138	111	144	295	371	518	529	440	571	184	318	543	600	407	353	110	
			60						225	165	213	279	407	437	249	567	619	73	131	181	217	371	187	139	
			40						378	183	159	142	130	103	156	614	682	144	68	68	65	100	121	434	

61 9 11 STA 20 6300

TIME 2225

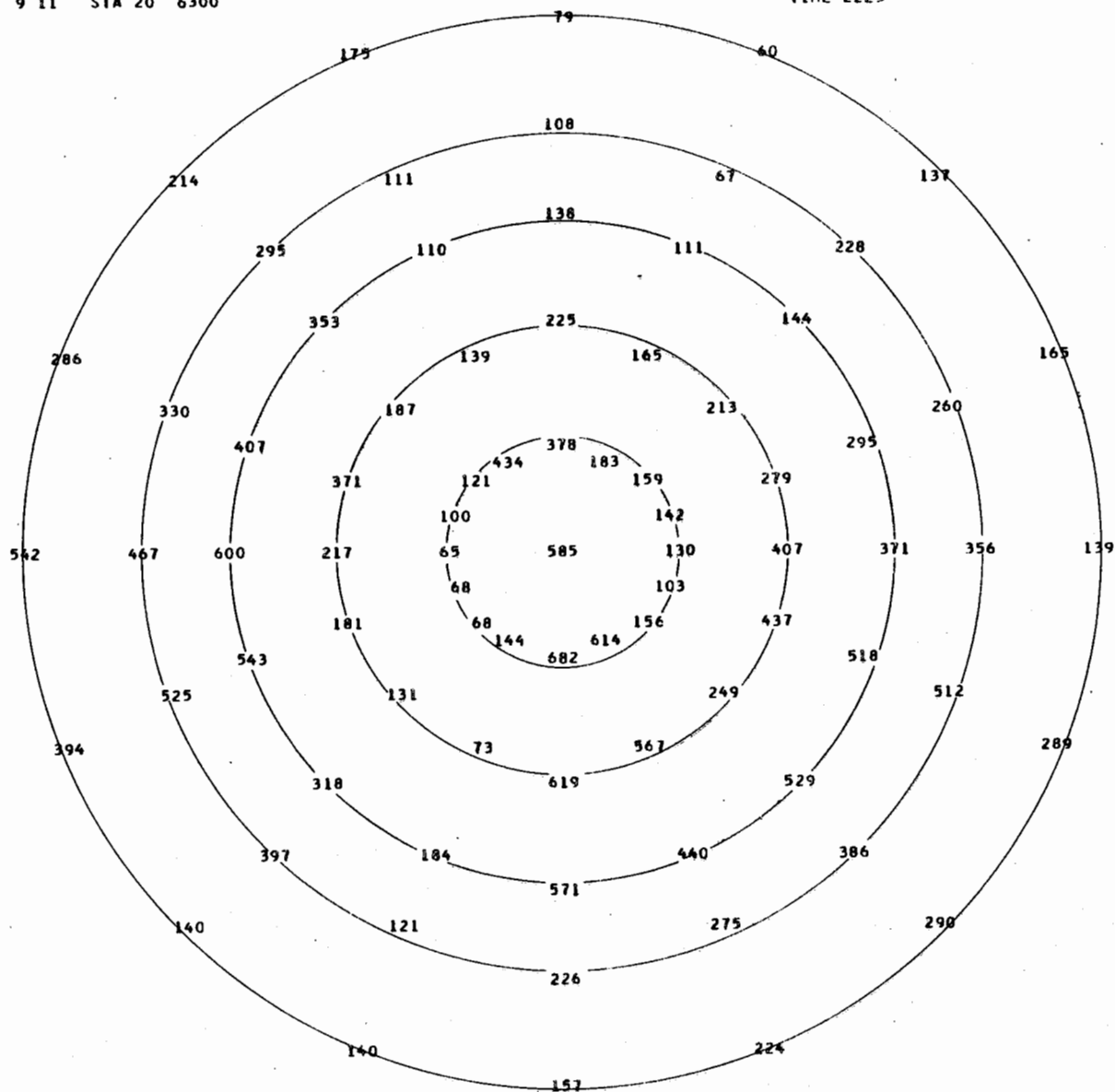


Figure 5

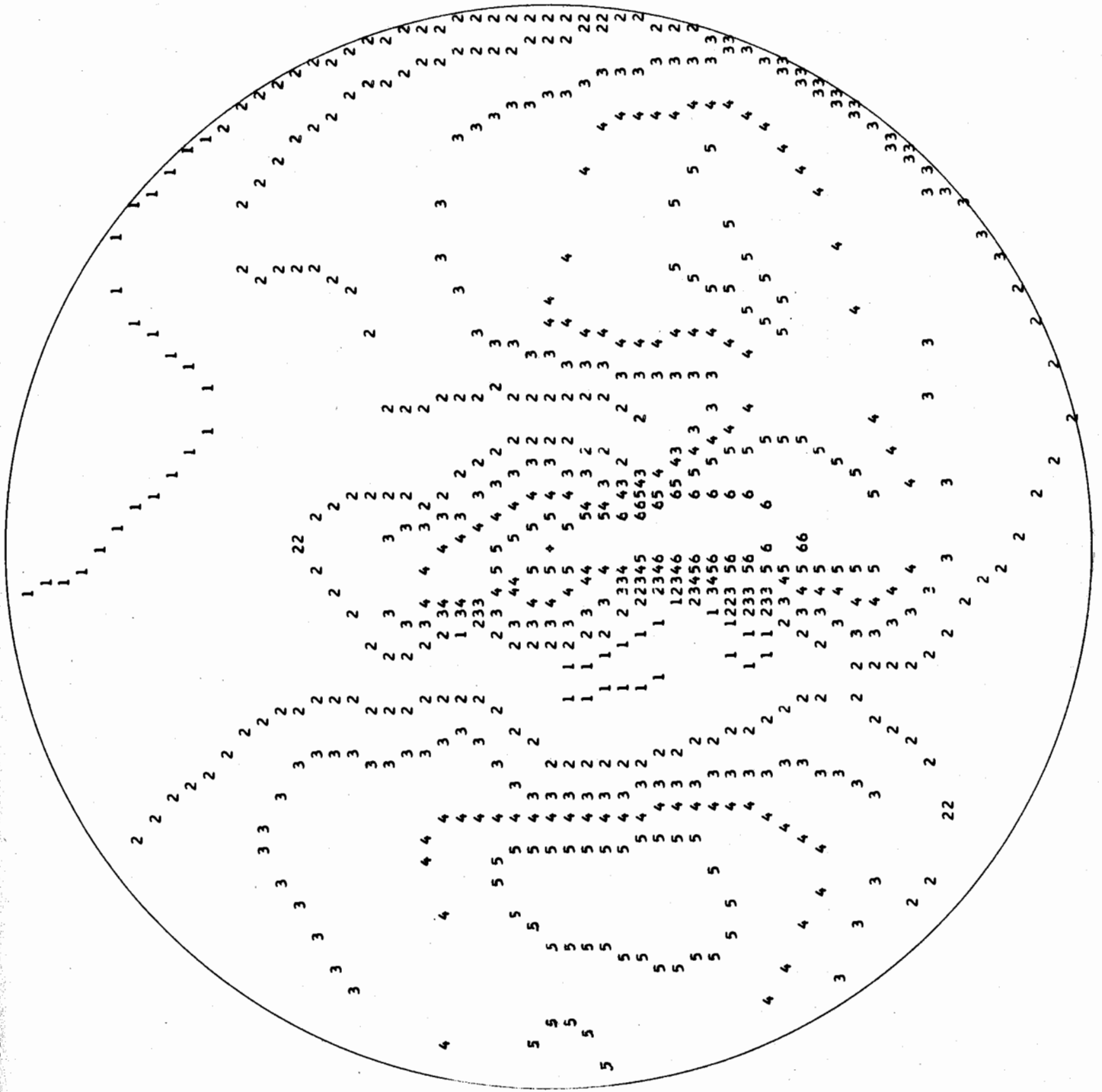


Figure 6

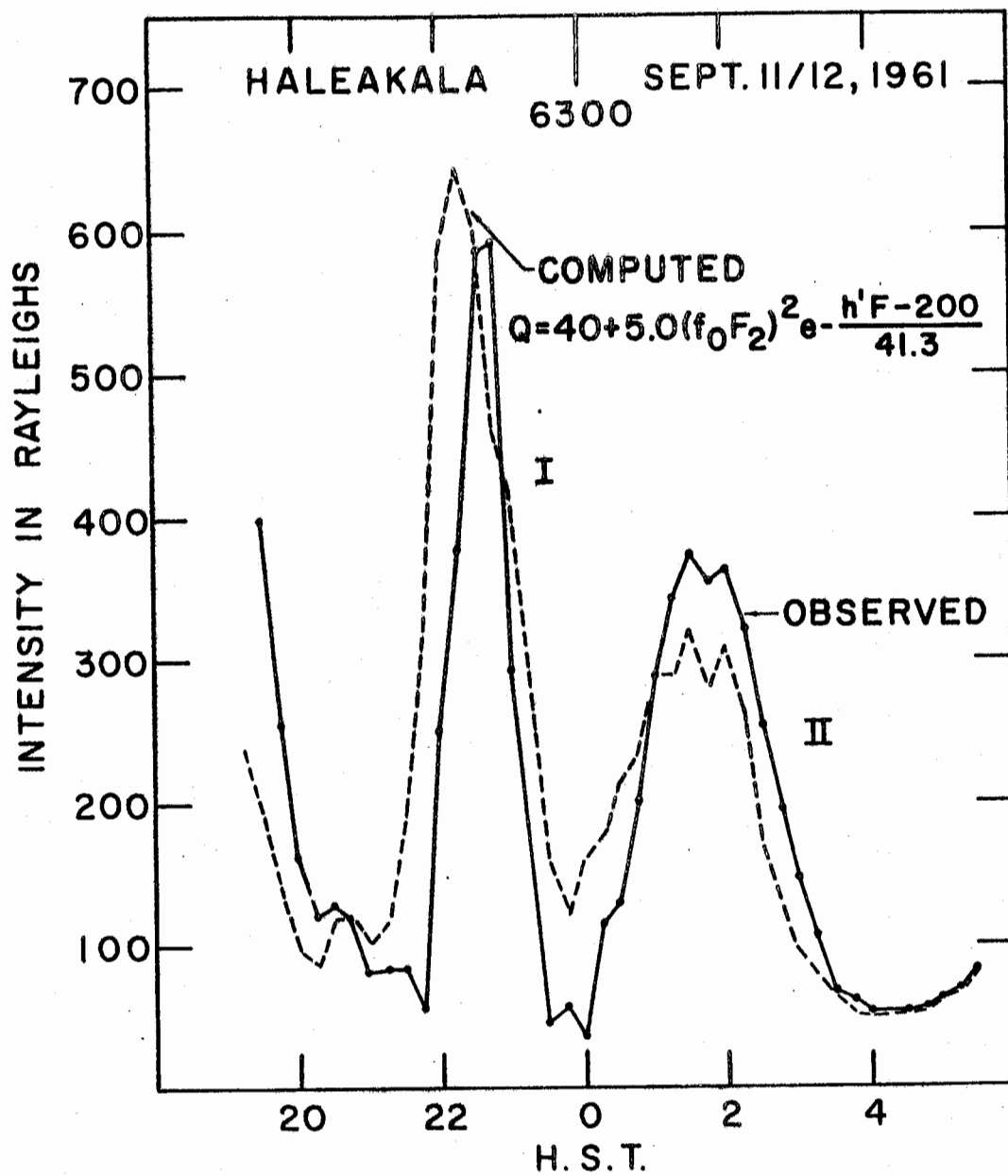


Figure 7

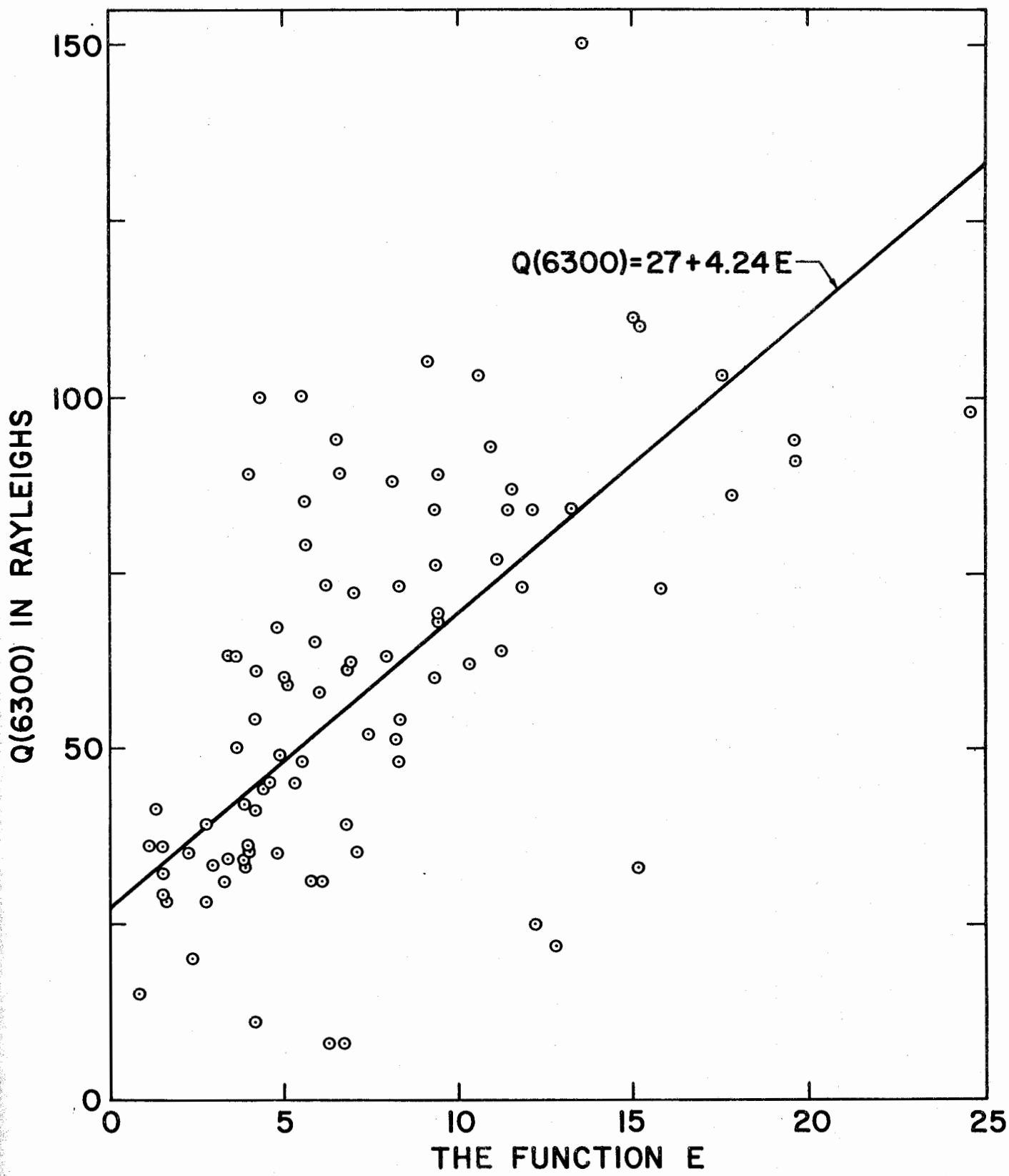


Figure 8

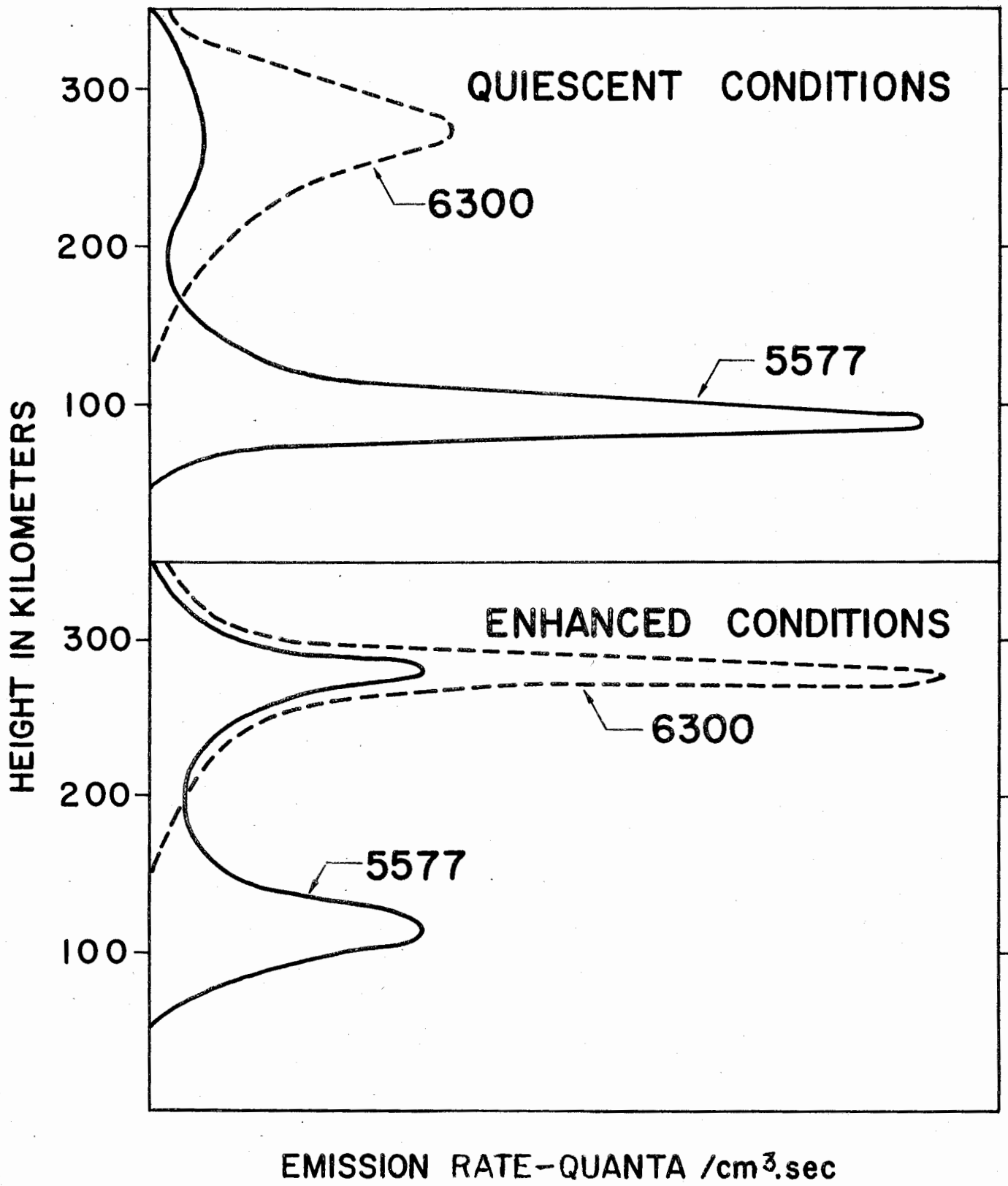


Figure 9

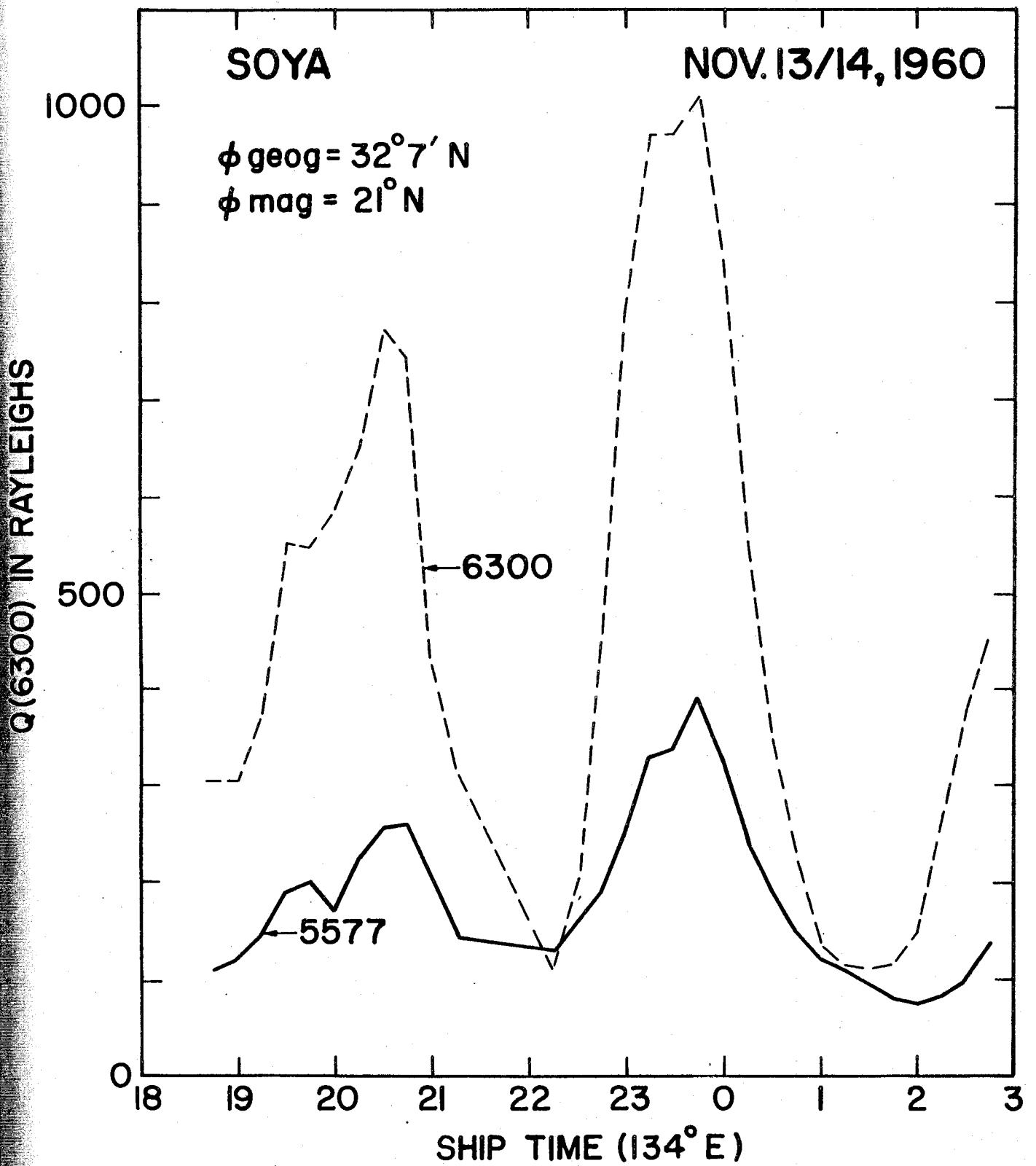


Figure 10

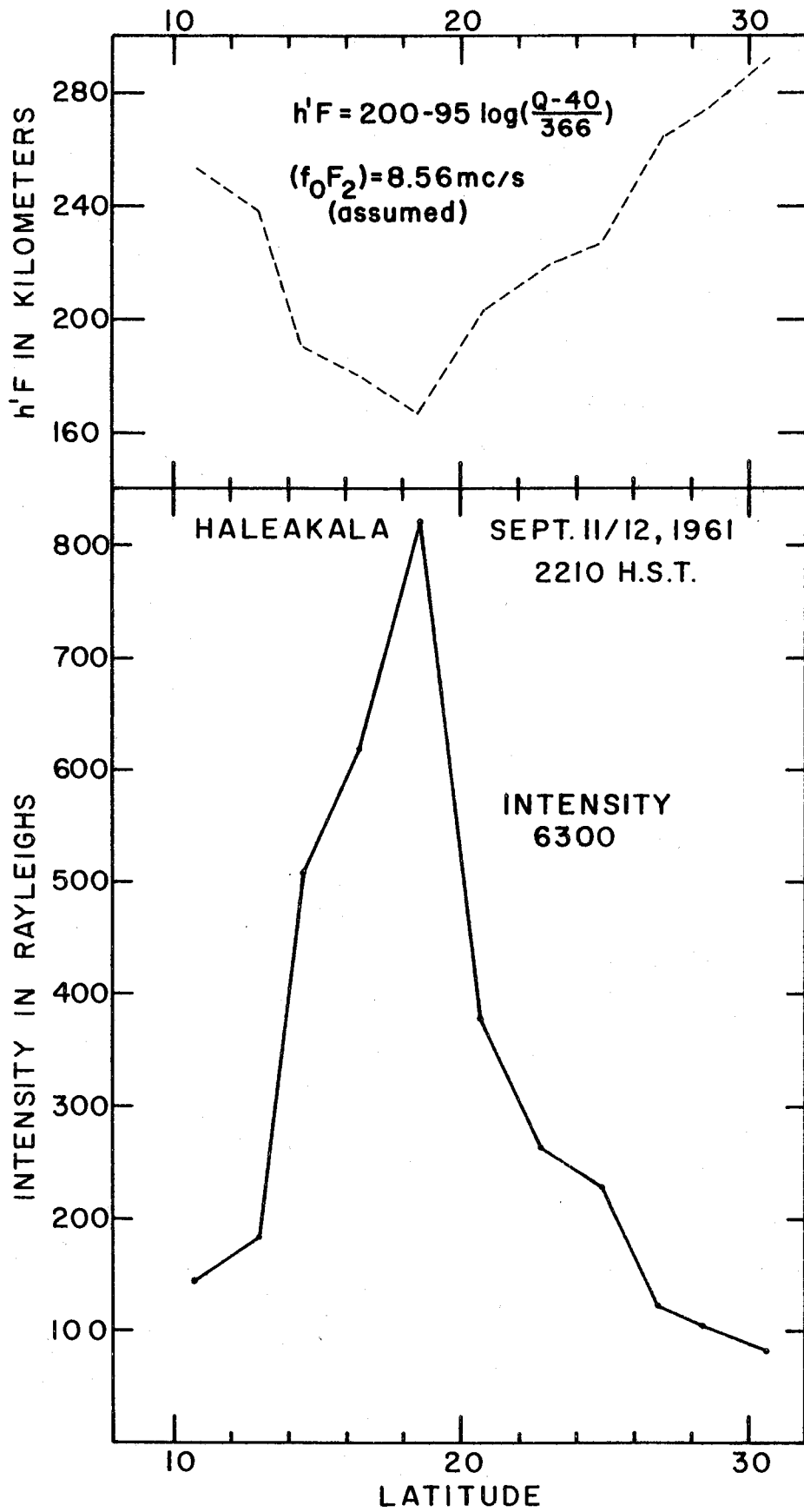


Figure 11.

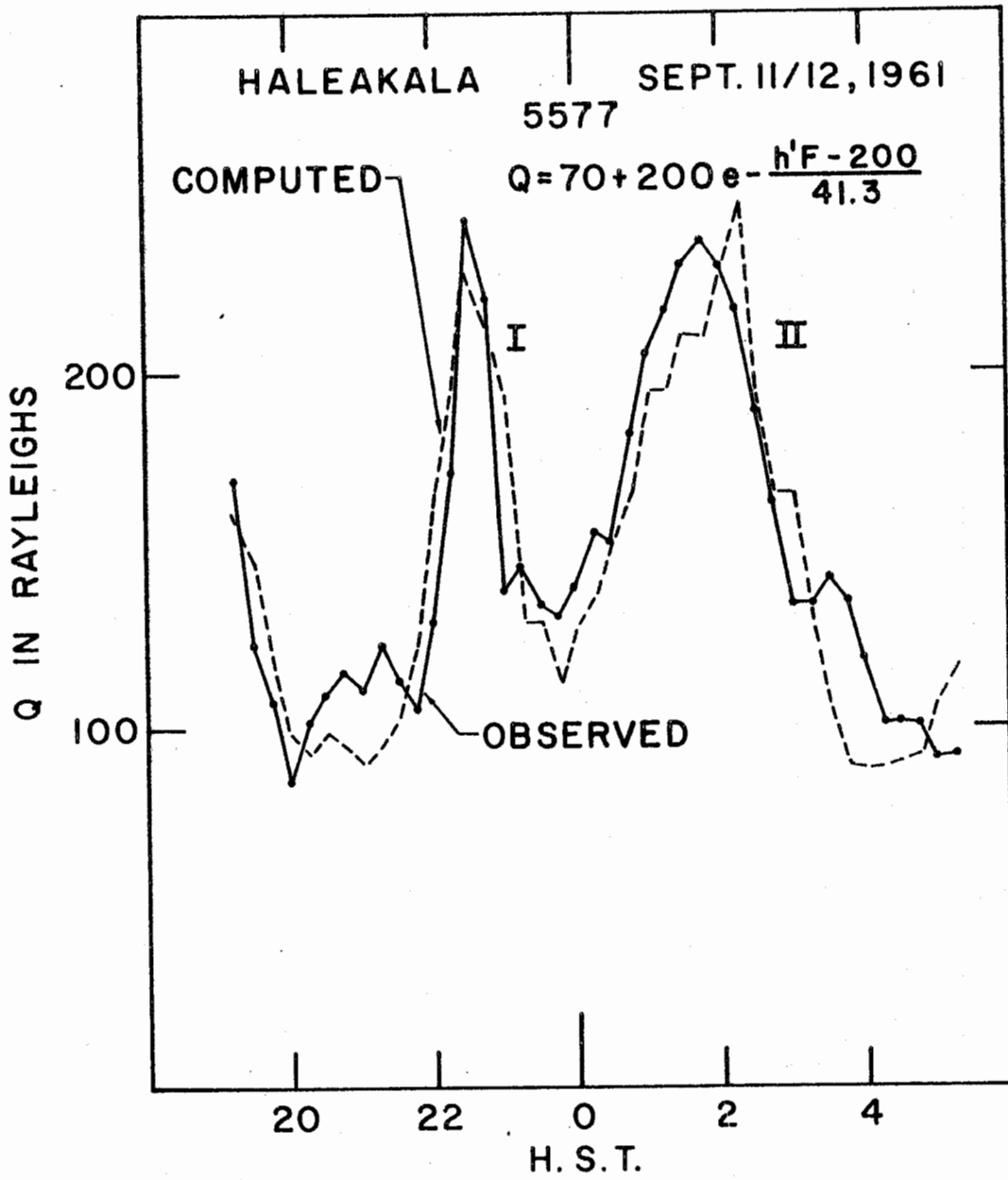


Figure 12

61/09/11
22:25
h'F

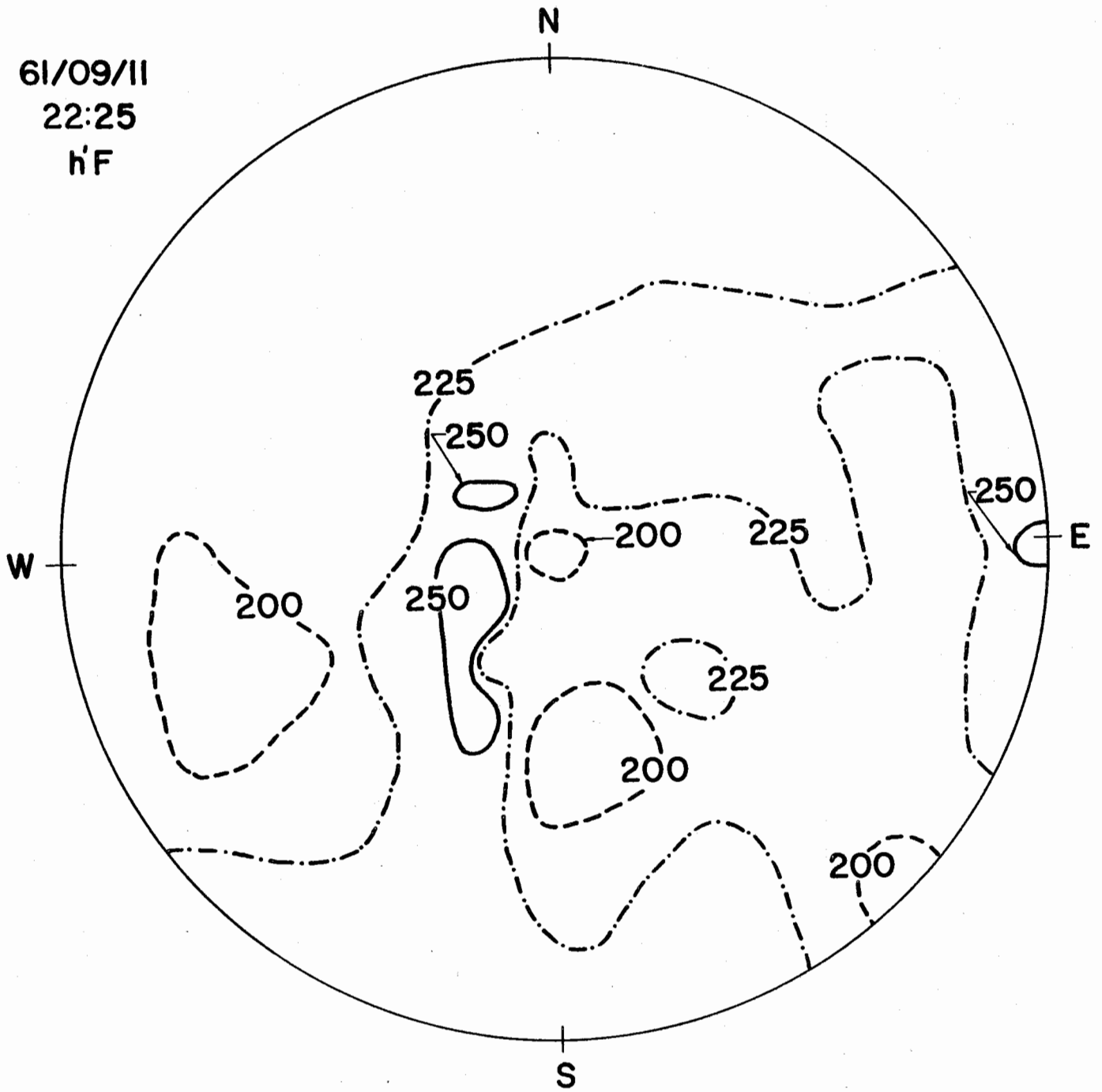


Figure 13A

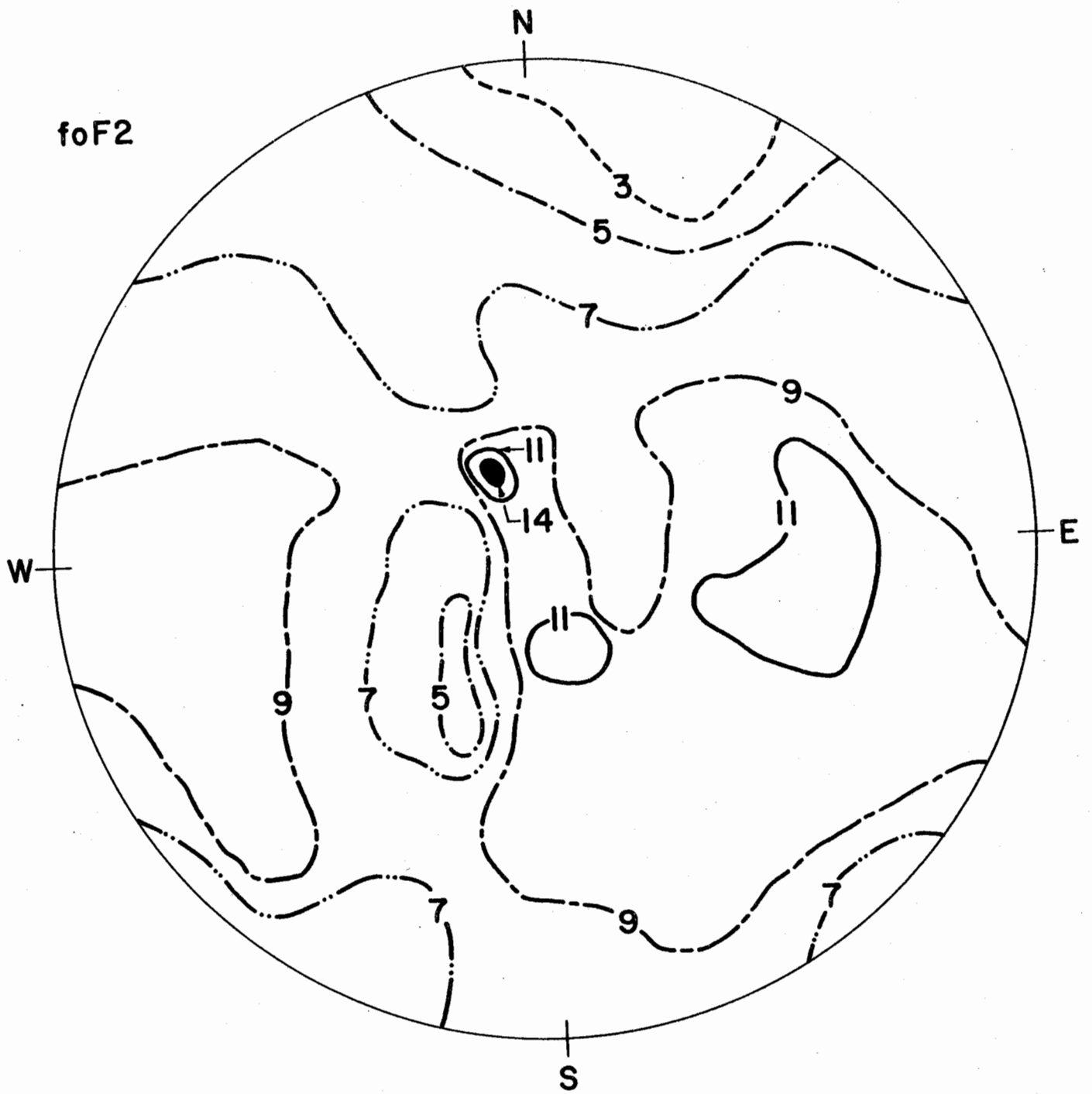


Figure 13B

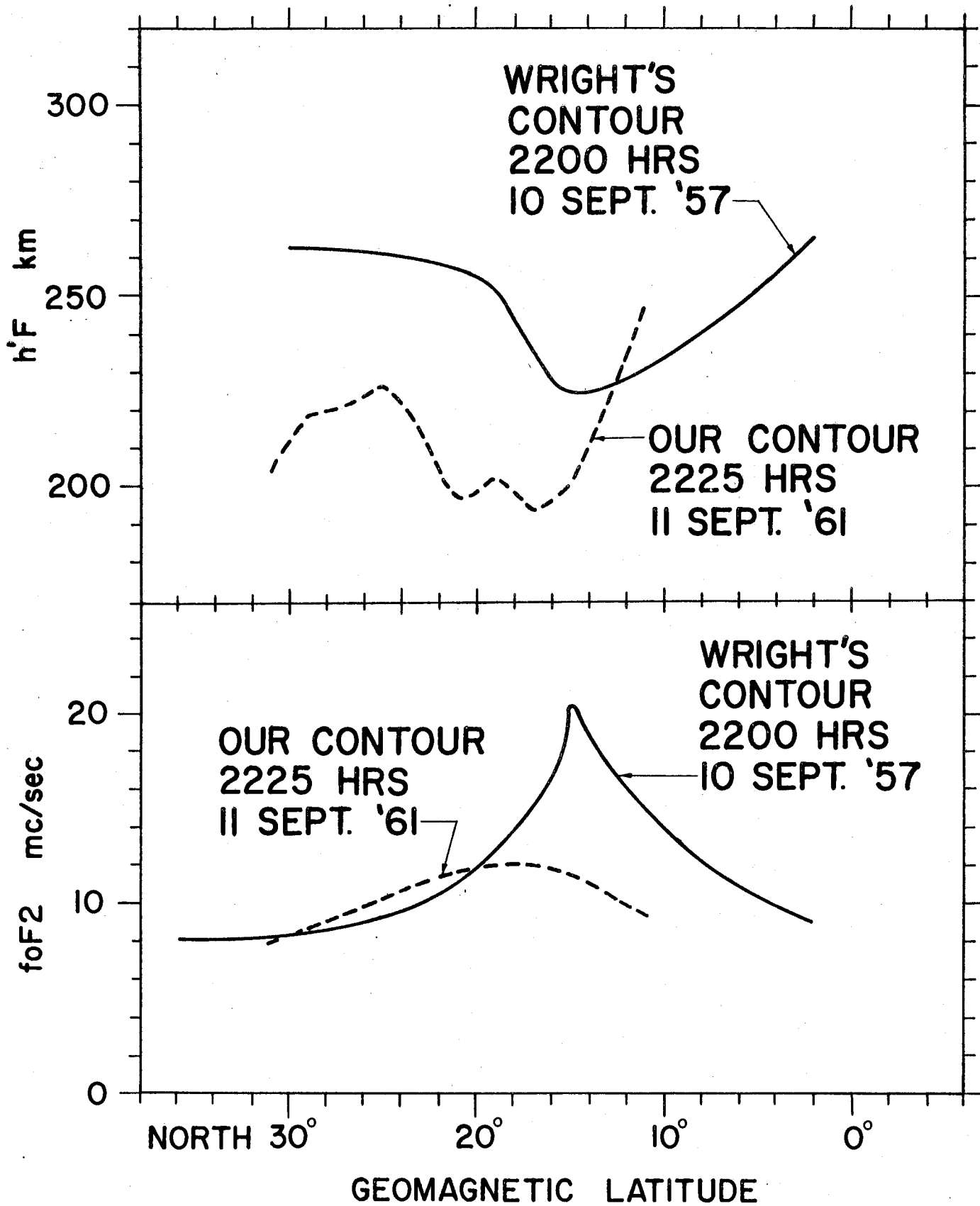


Figure 14

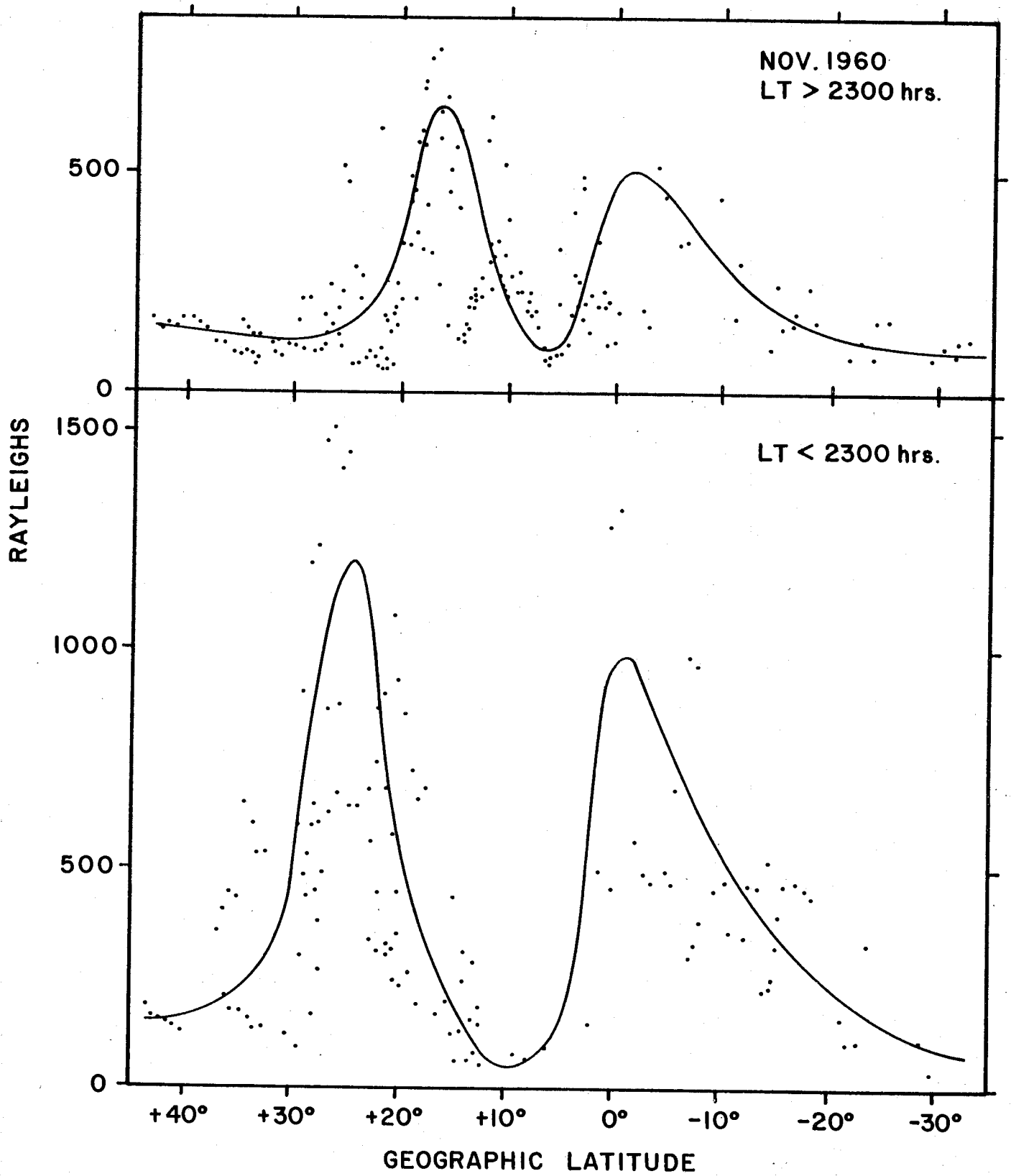


Figure 15

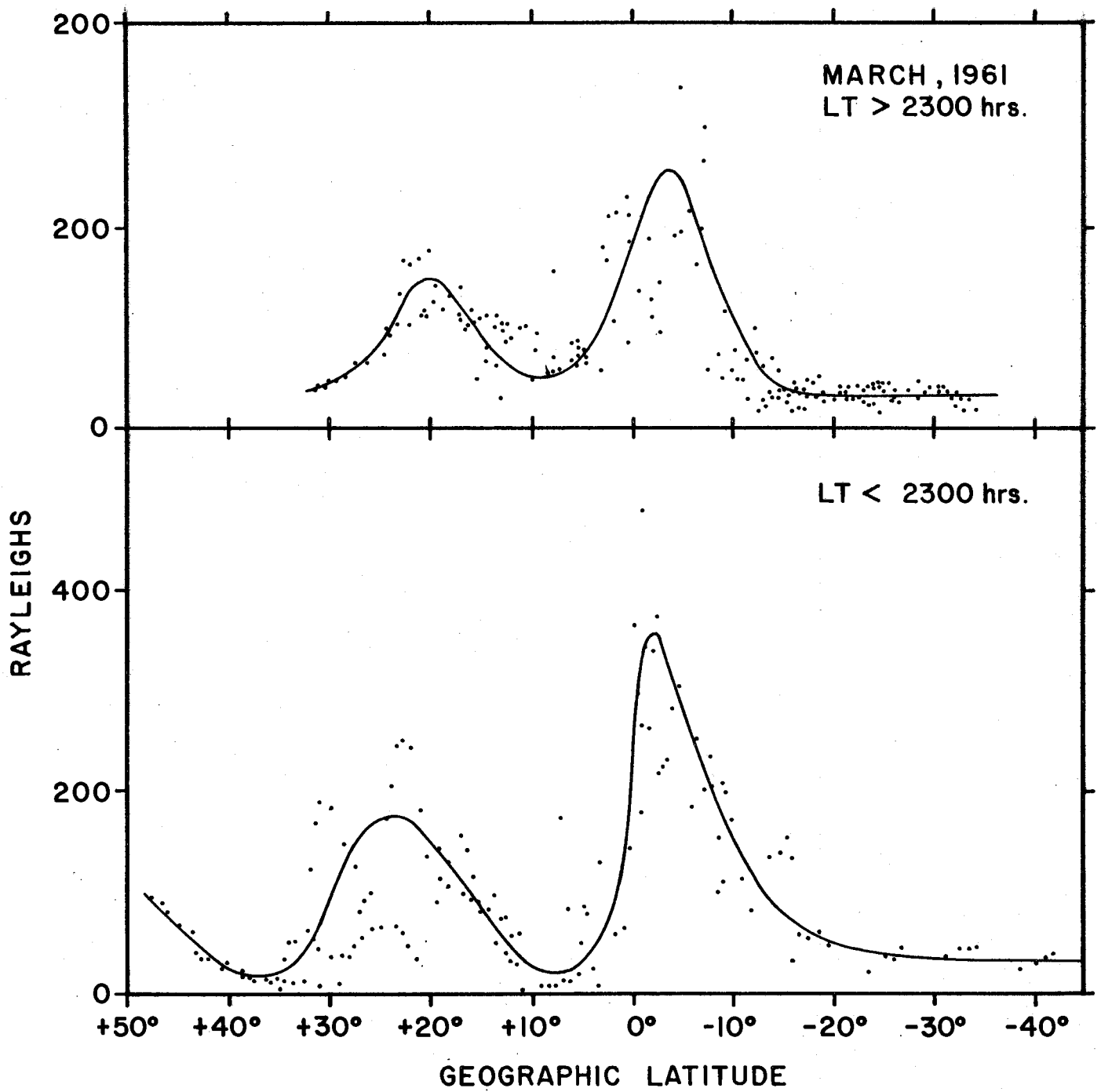


Figure 16

ELTANIN EXPEDITION
MEAN Q (6300) OF ONE
TRIP, MARCH-NOV., 1963

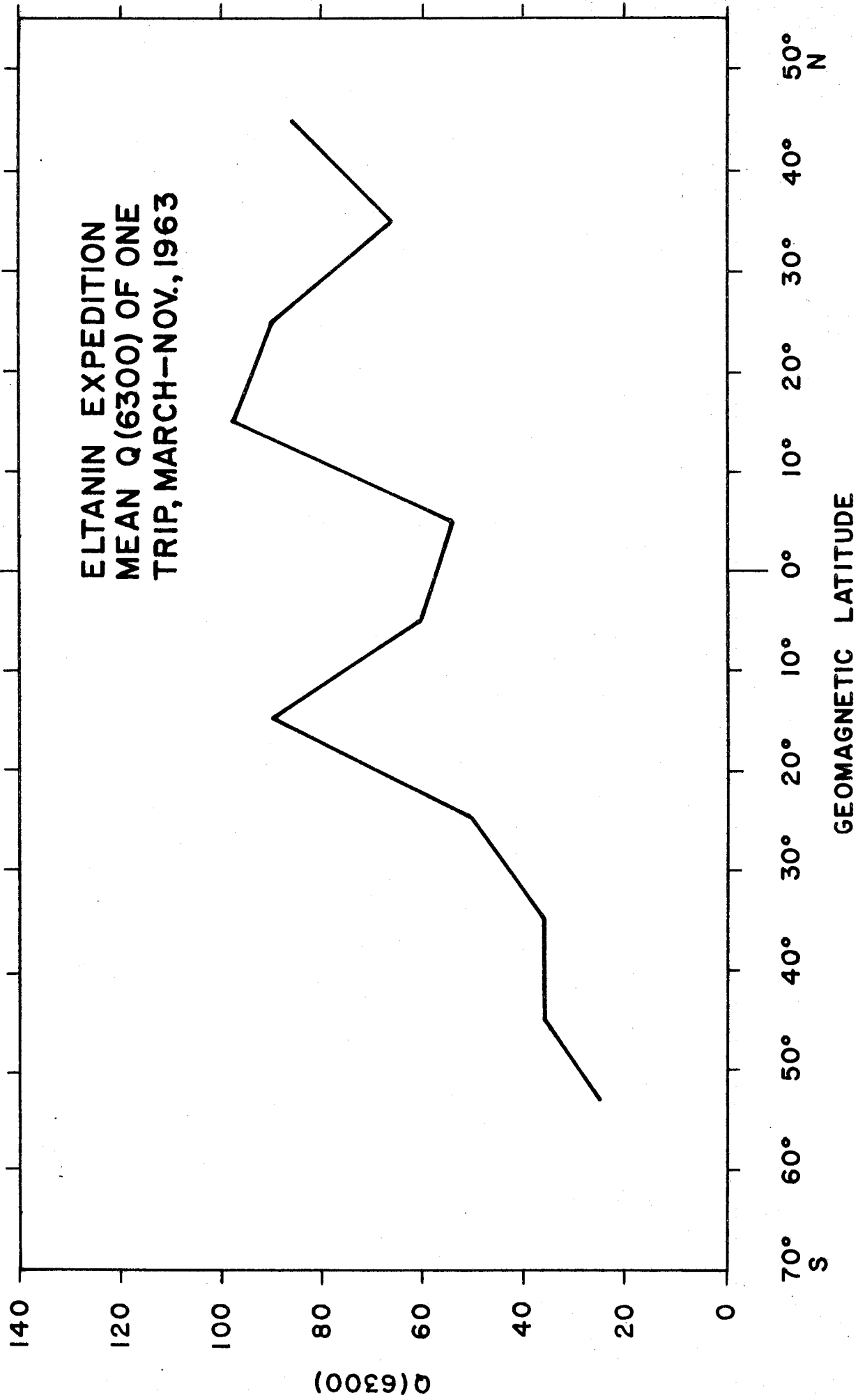


Figure 17

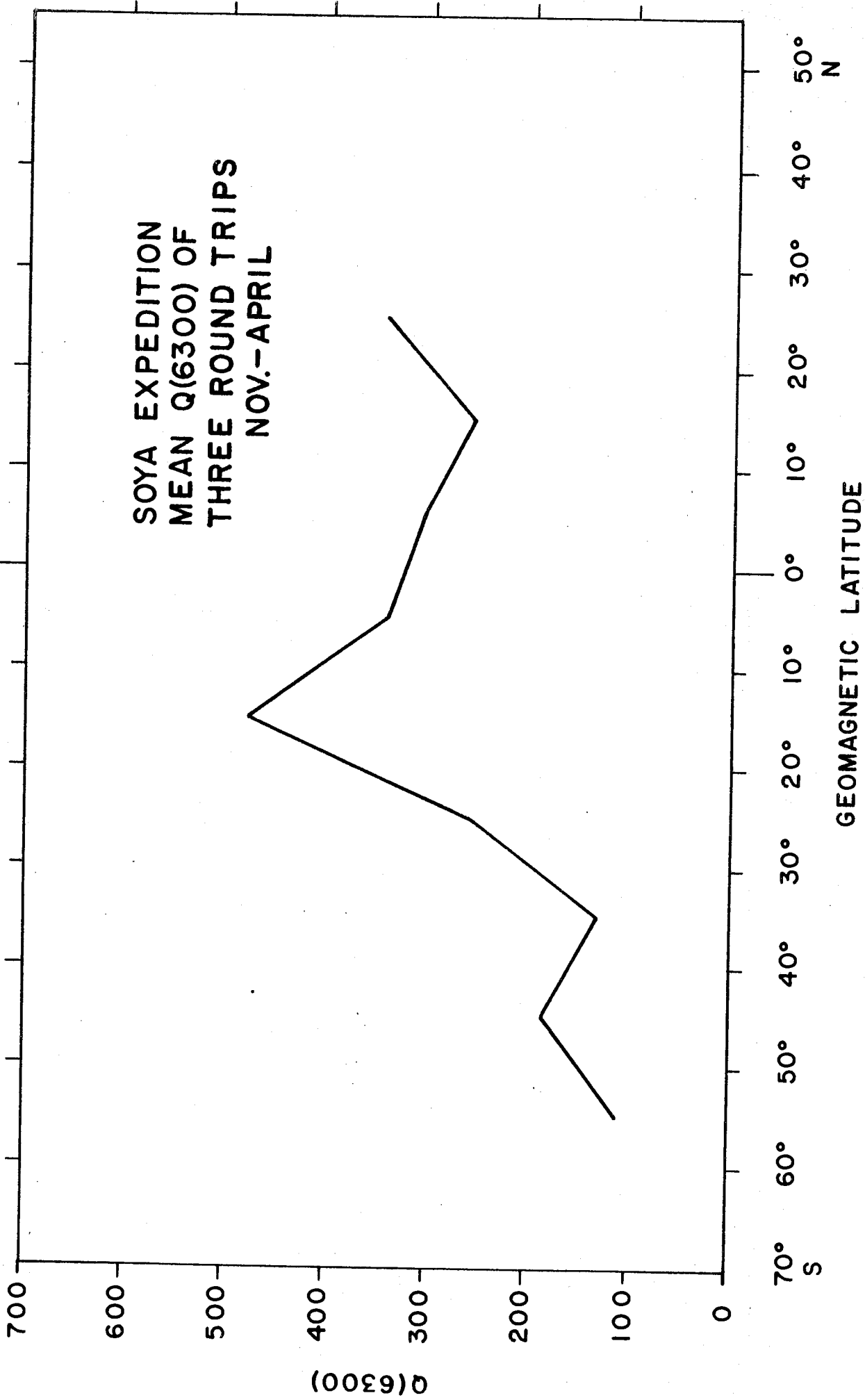


Figure 18

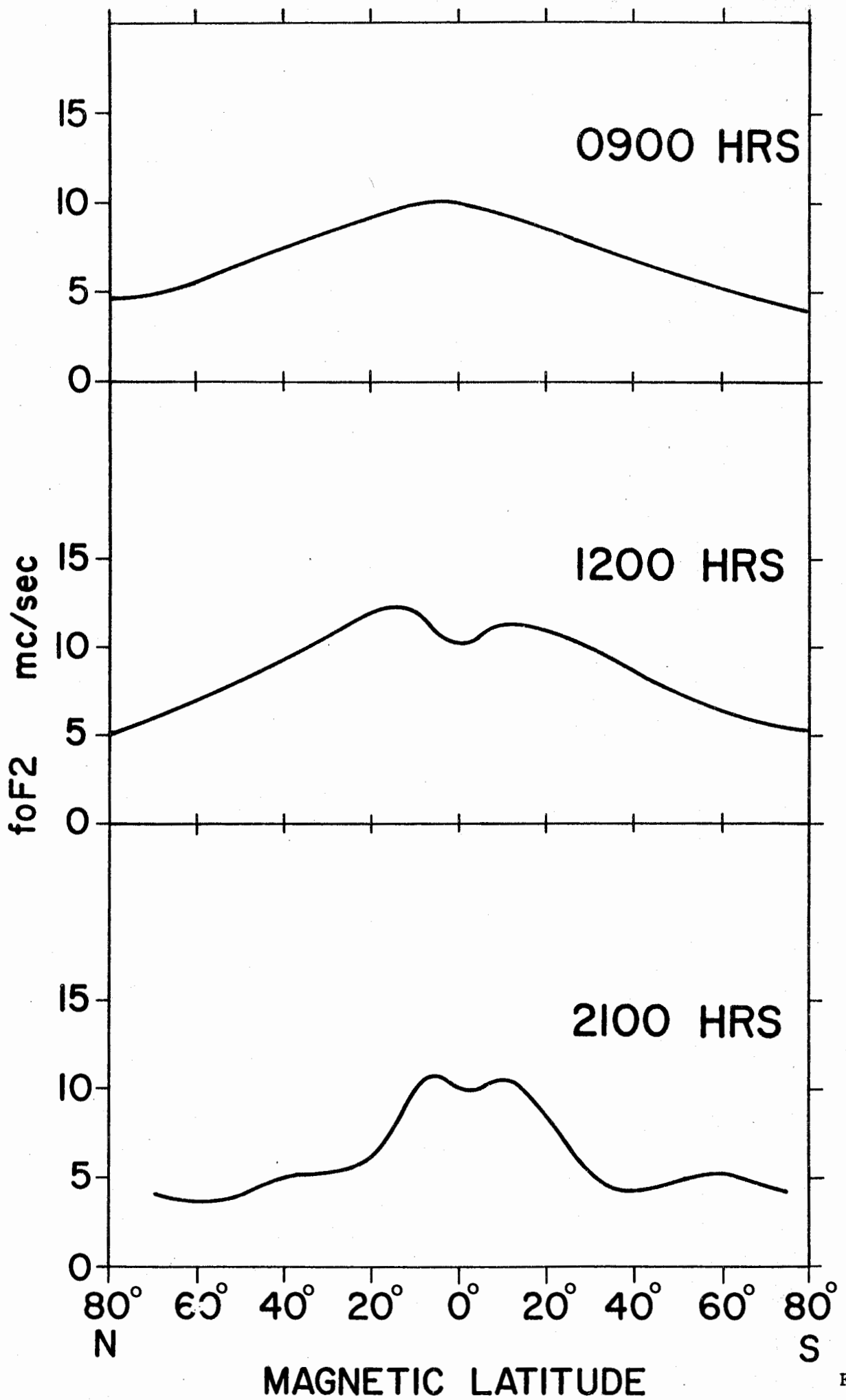


Figure 19

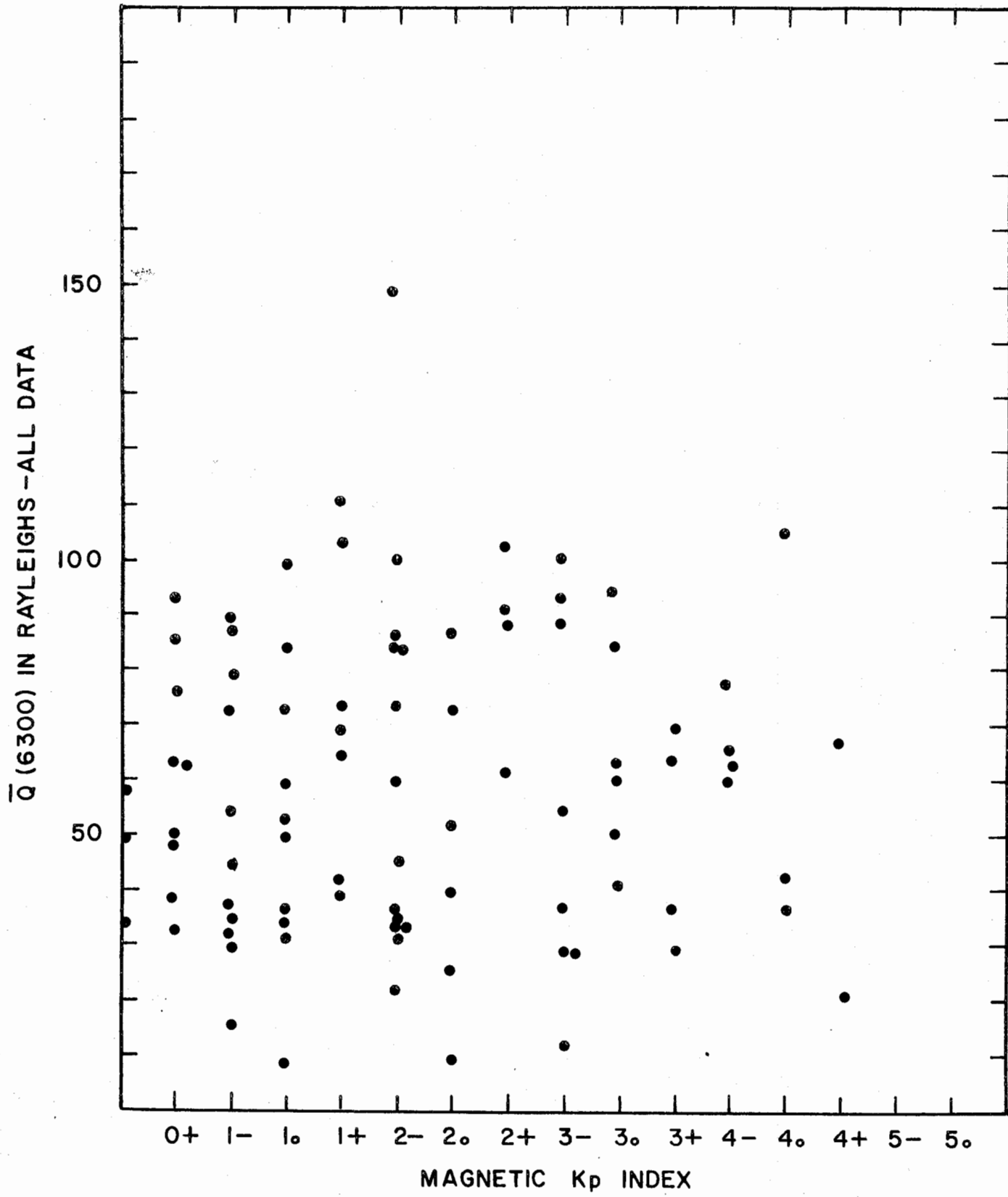


Figure 20

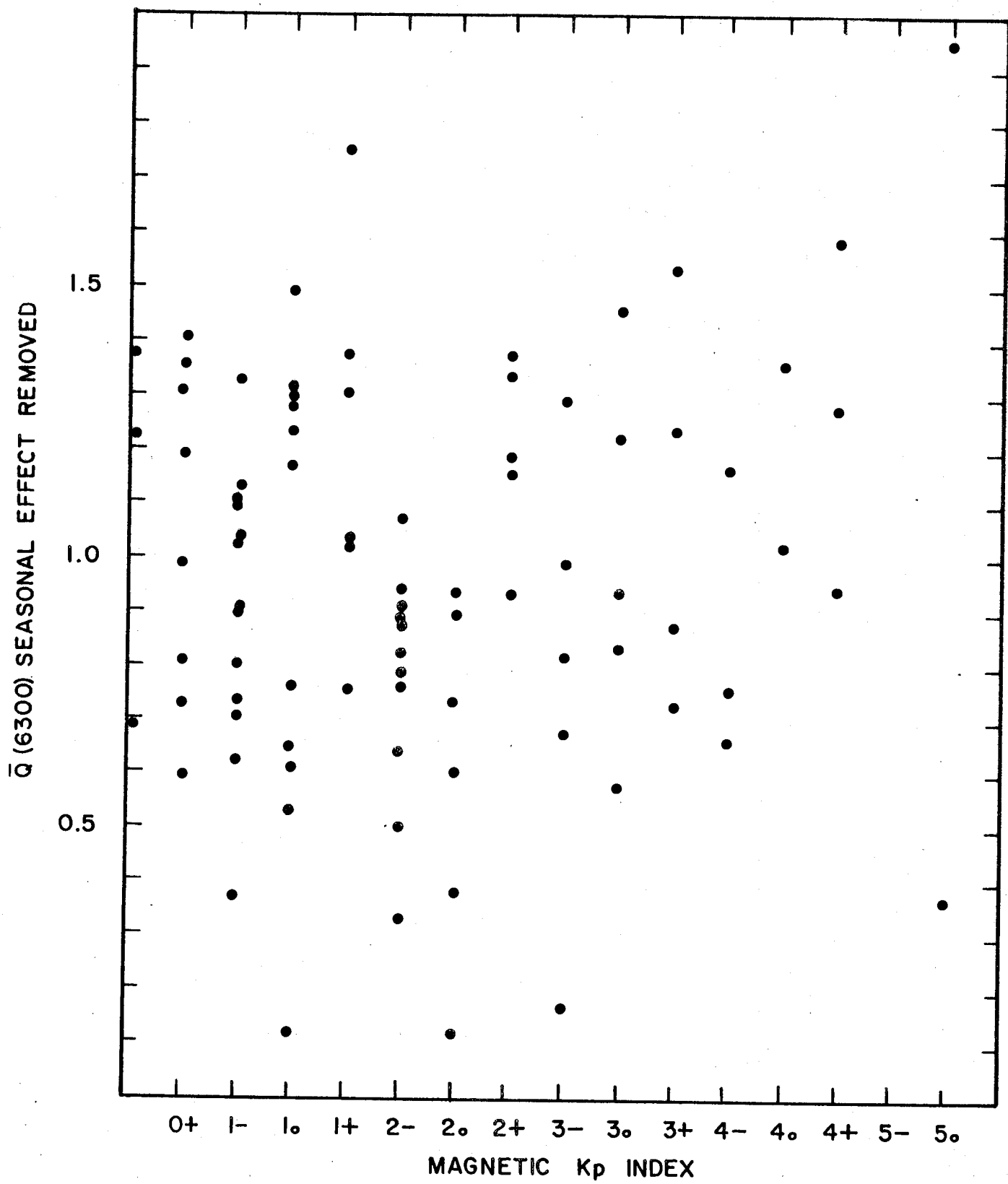


Figure 21

\bar{Q} (6300) IN RAYLEIGHS - SEPT., OCT., NOV., 1963

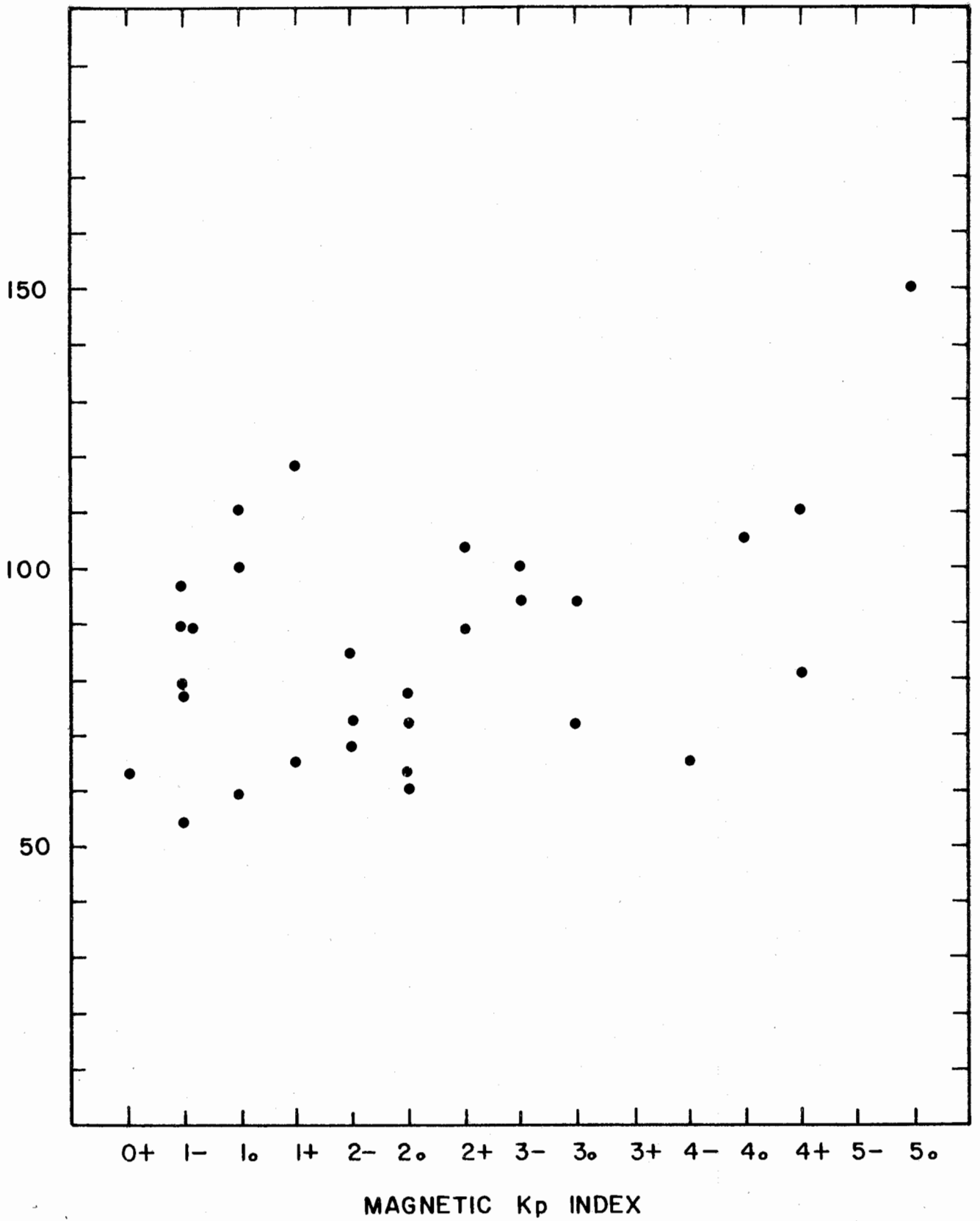
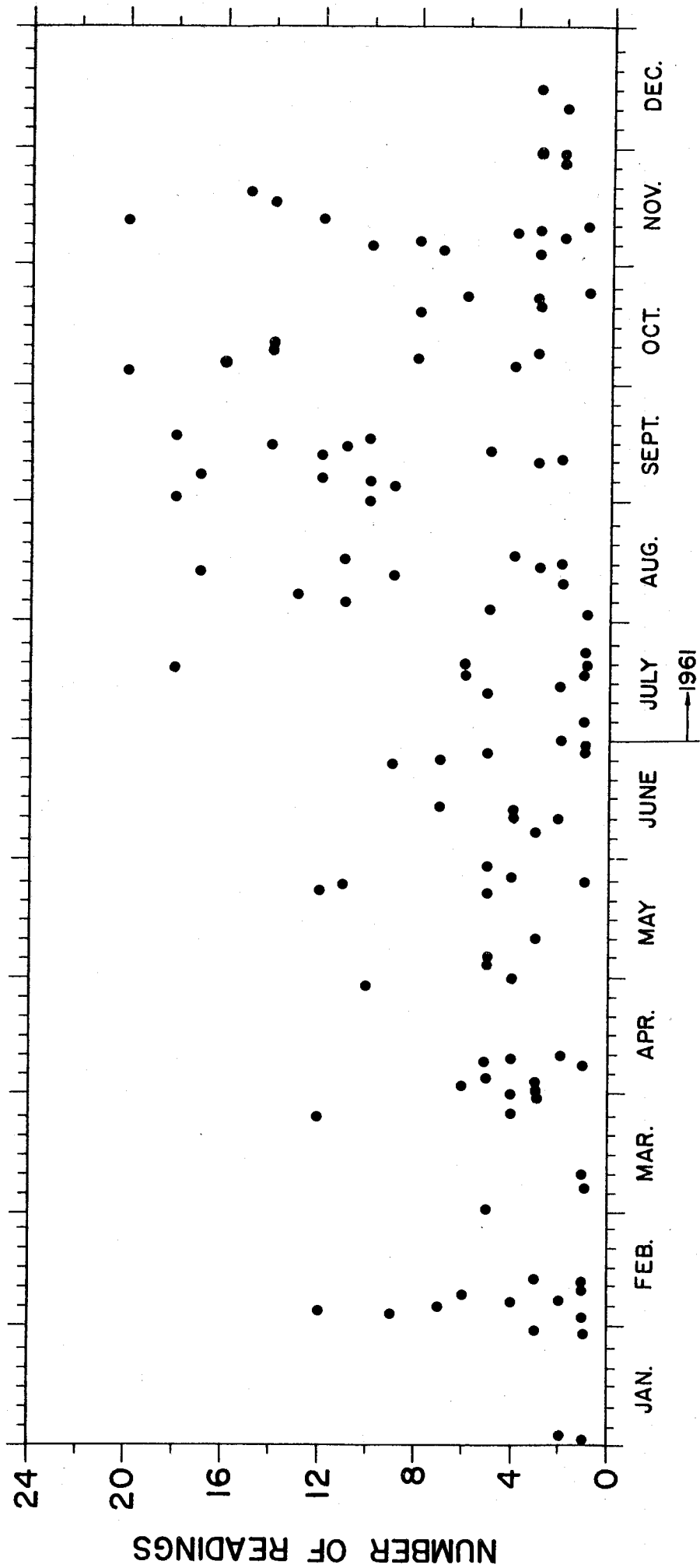
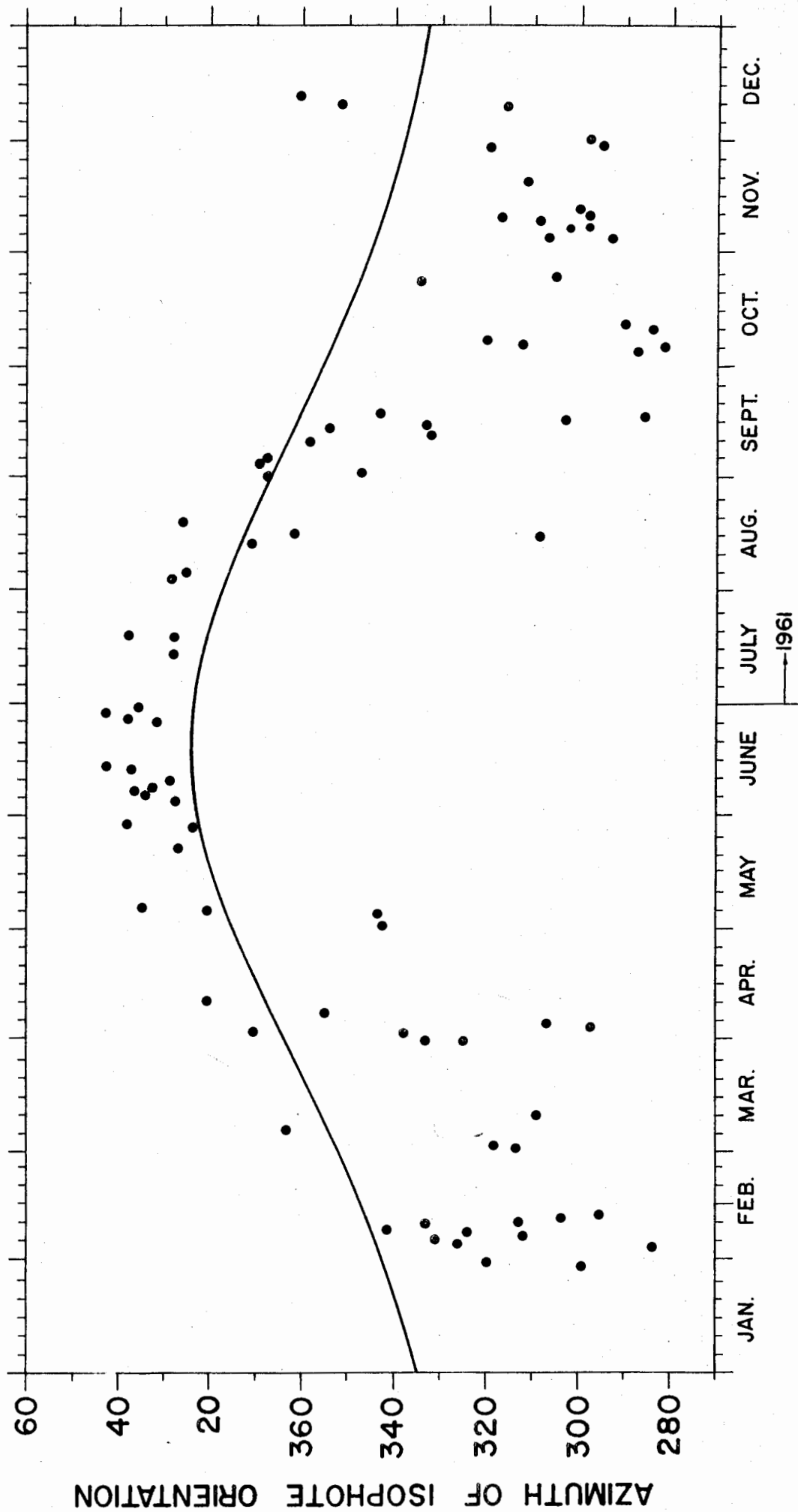


Figure 22



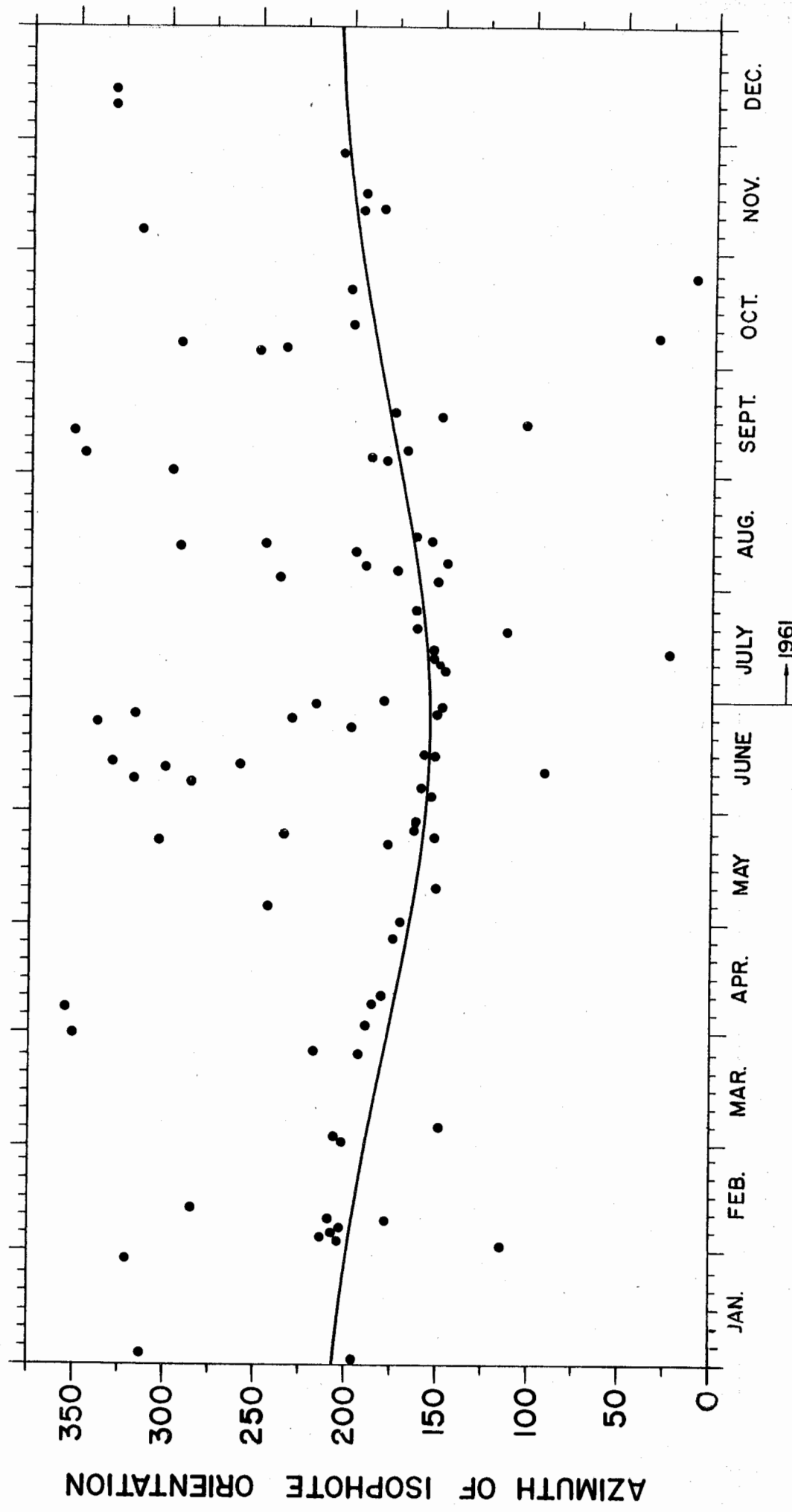
JULY 1961 TO JULY 1962

Figure 23



JULY 1961 TO JULY 1962

Figure 24



JULY 1961 TO JULY 1962

Figure 25

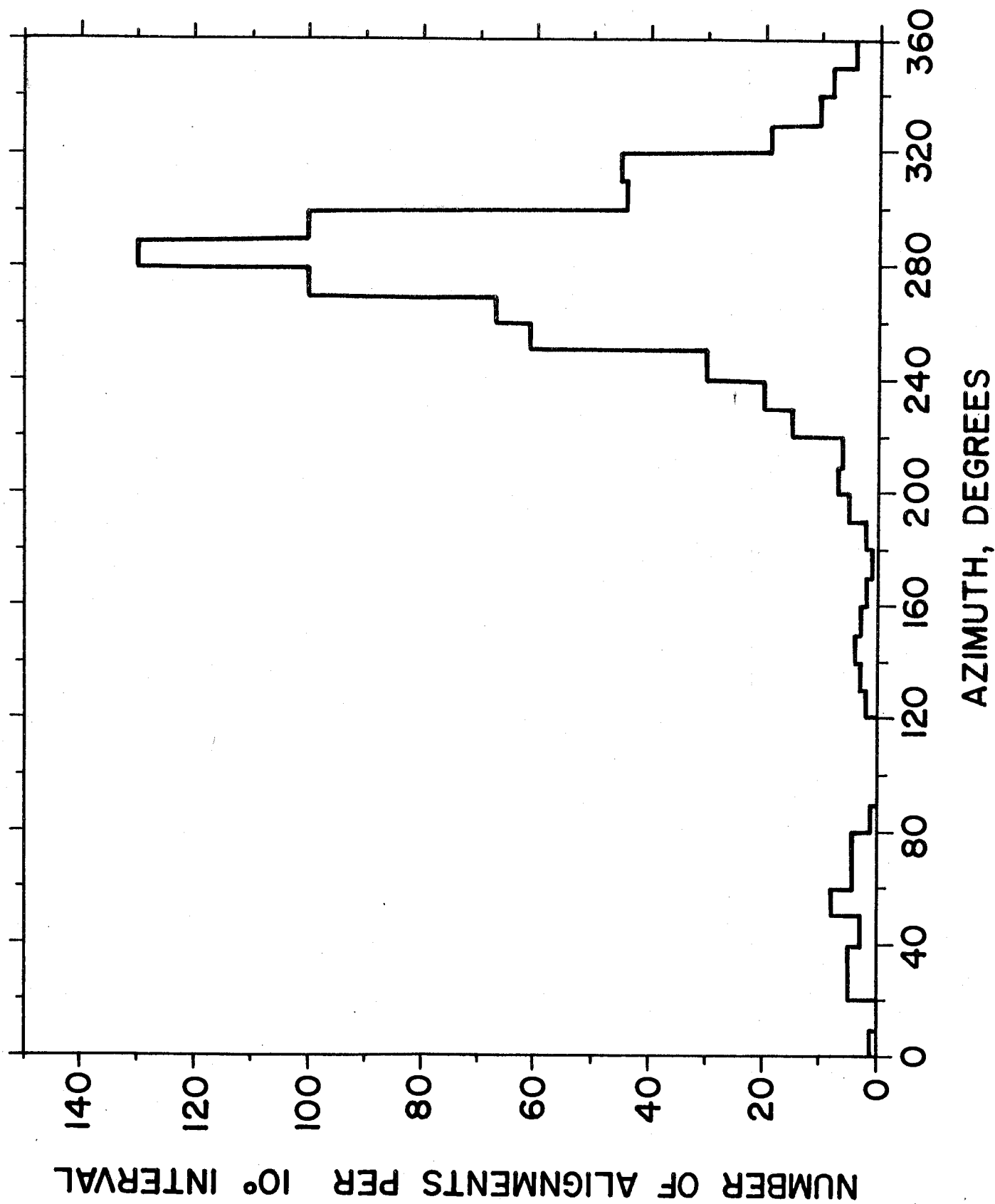


Figure 26

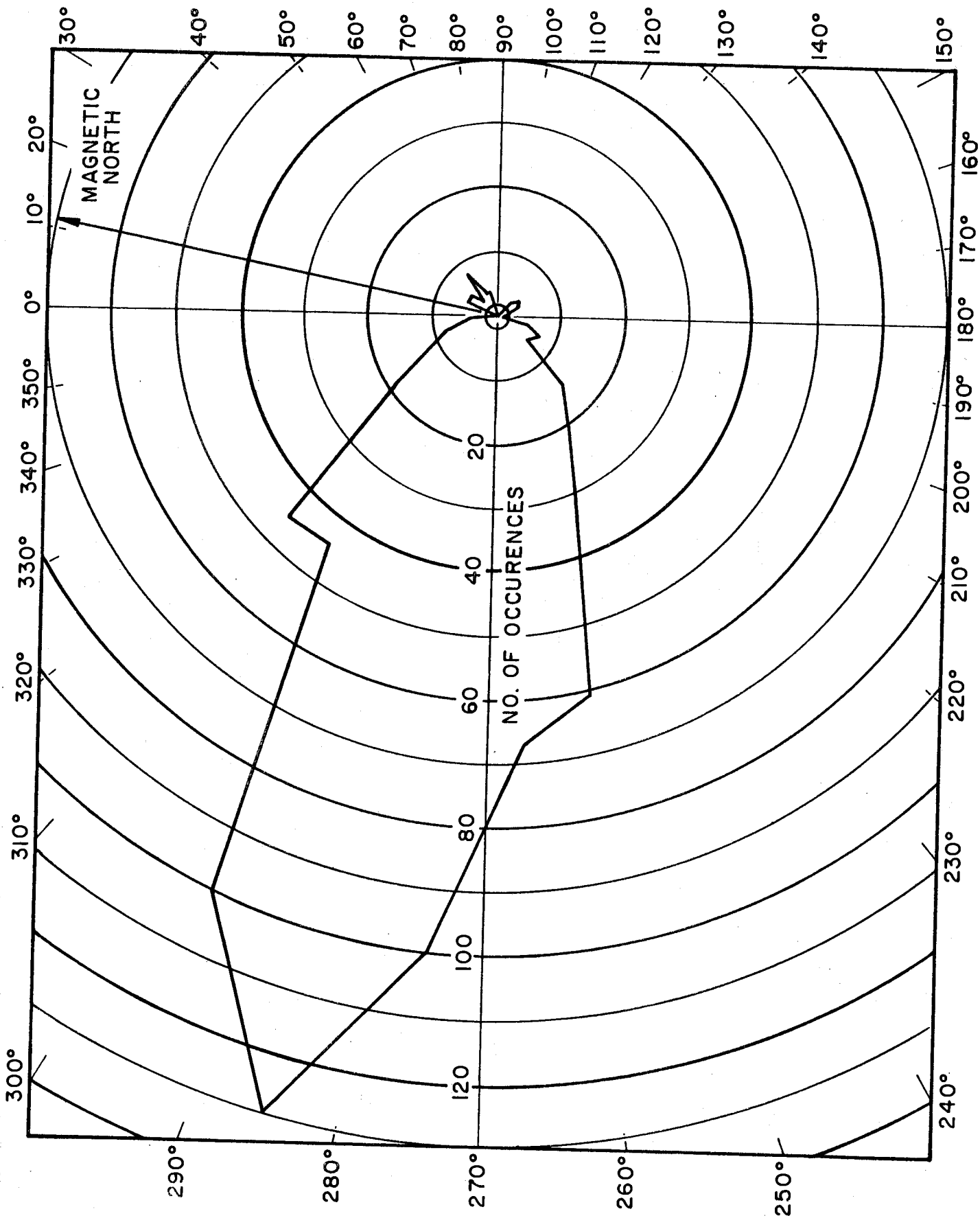


Figure 27

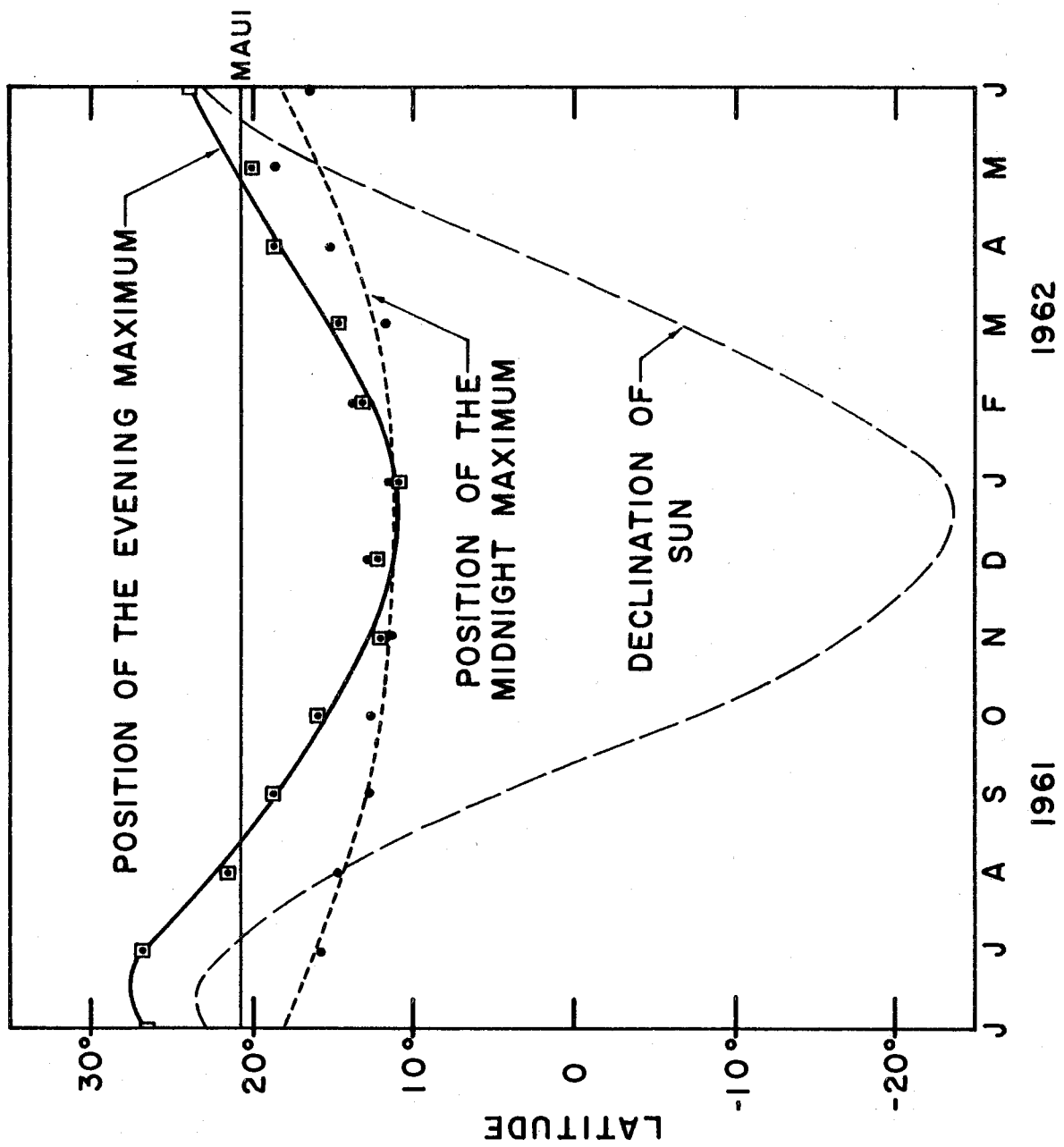


Figure 28

RELATIVE INTENSITY Q(6300)

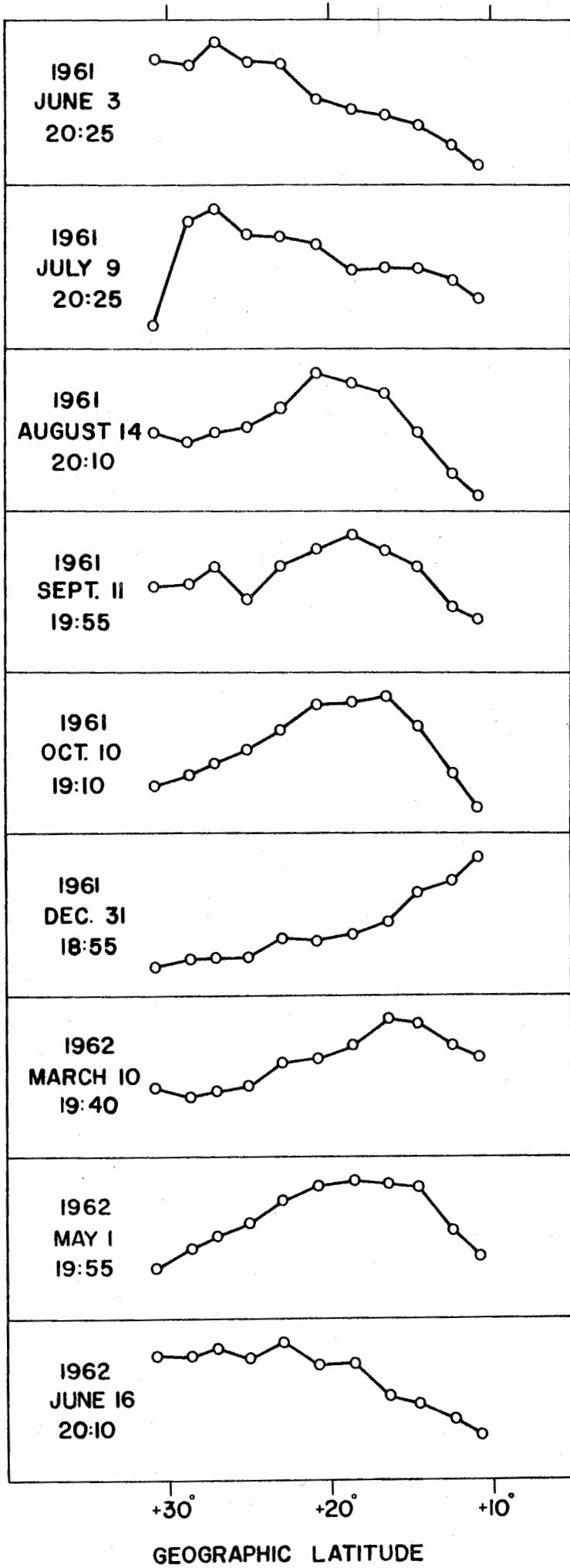


Figure 29

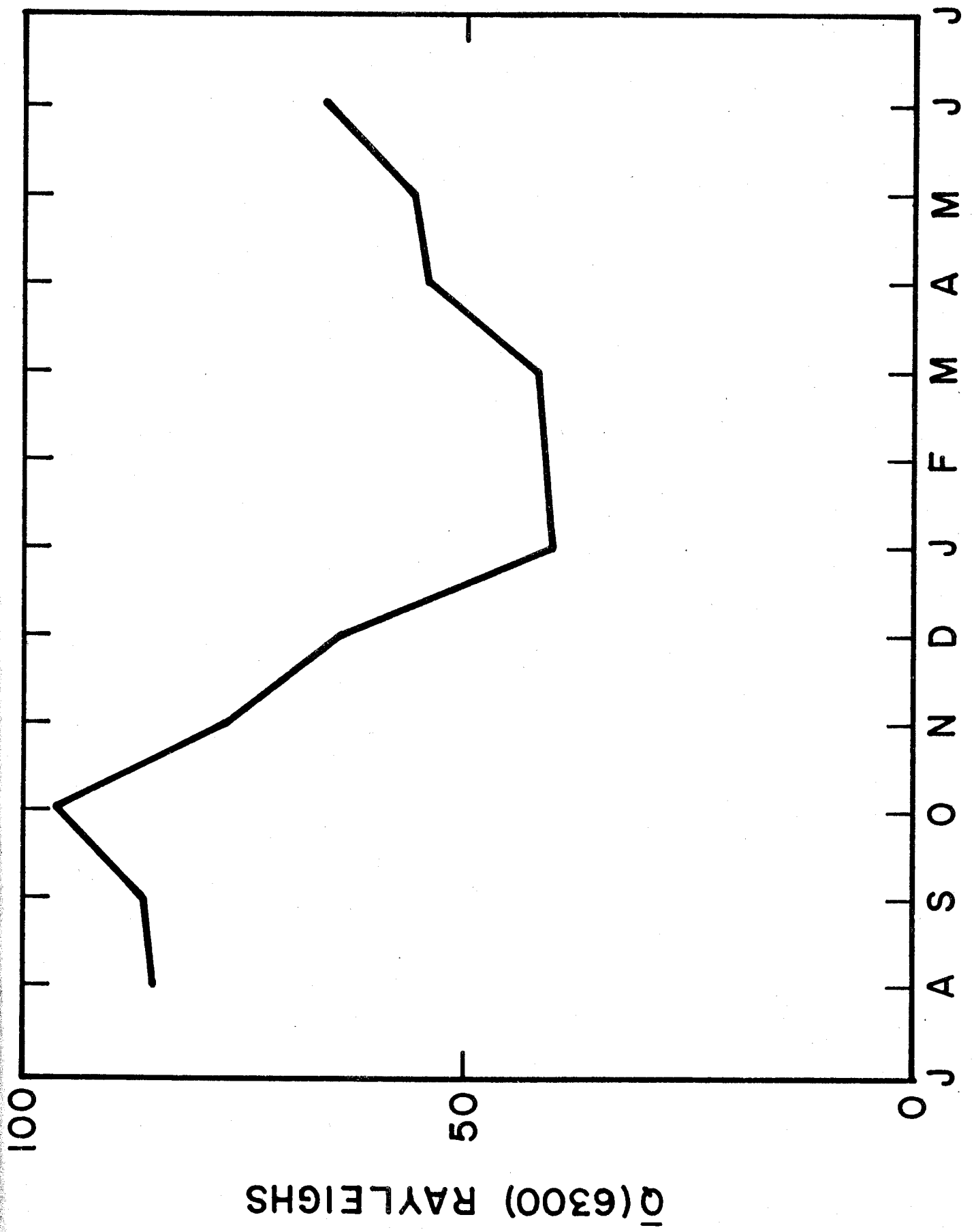


Figure 30

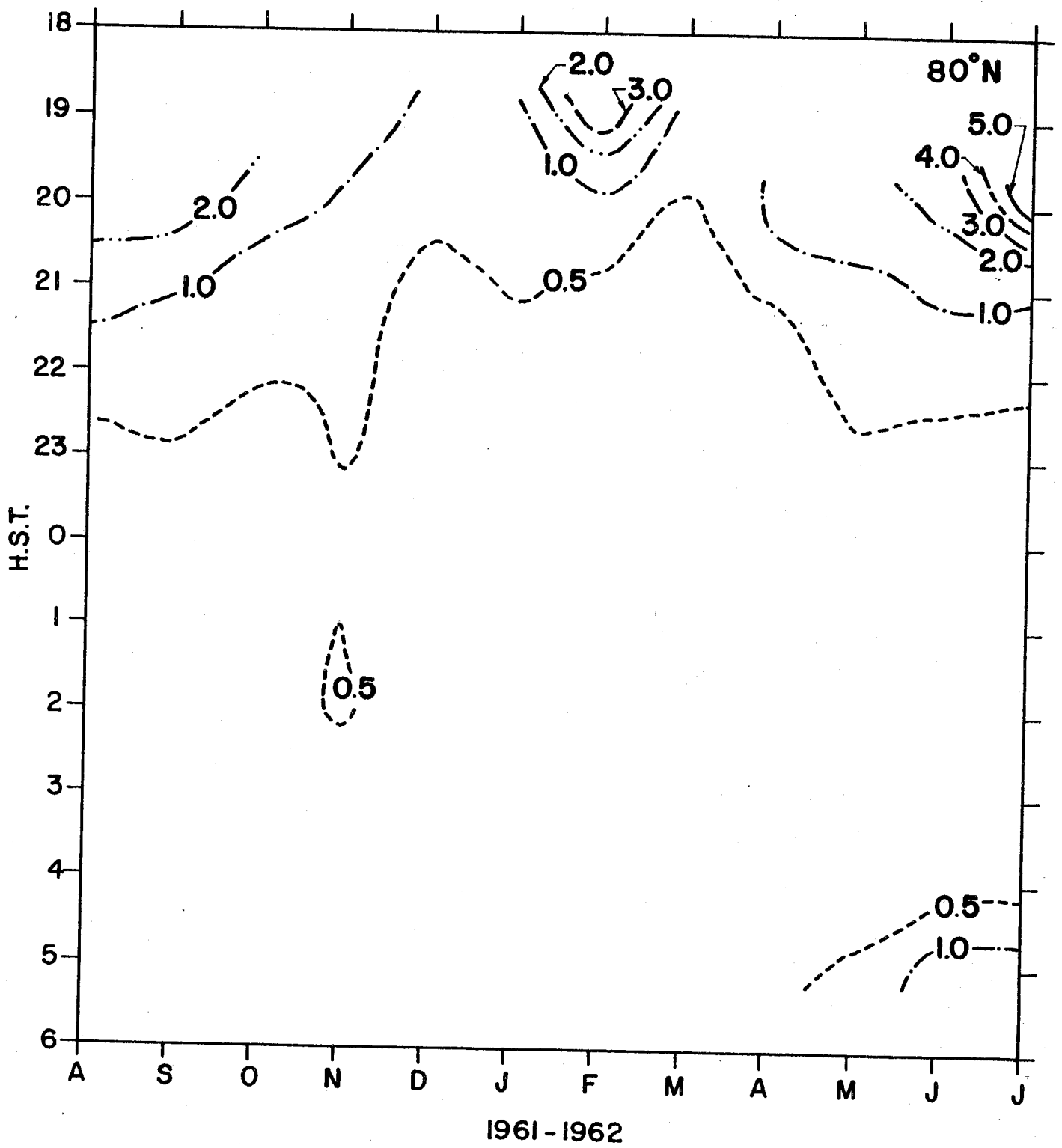


Figure 31

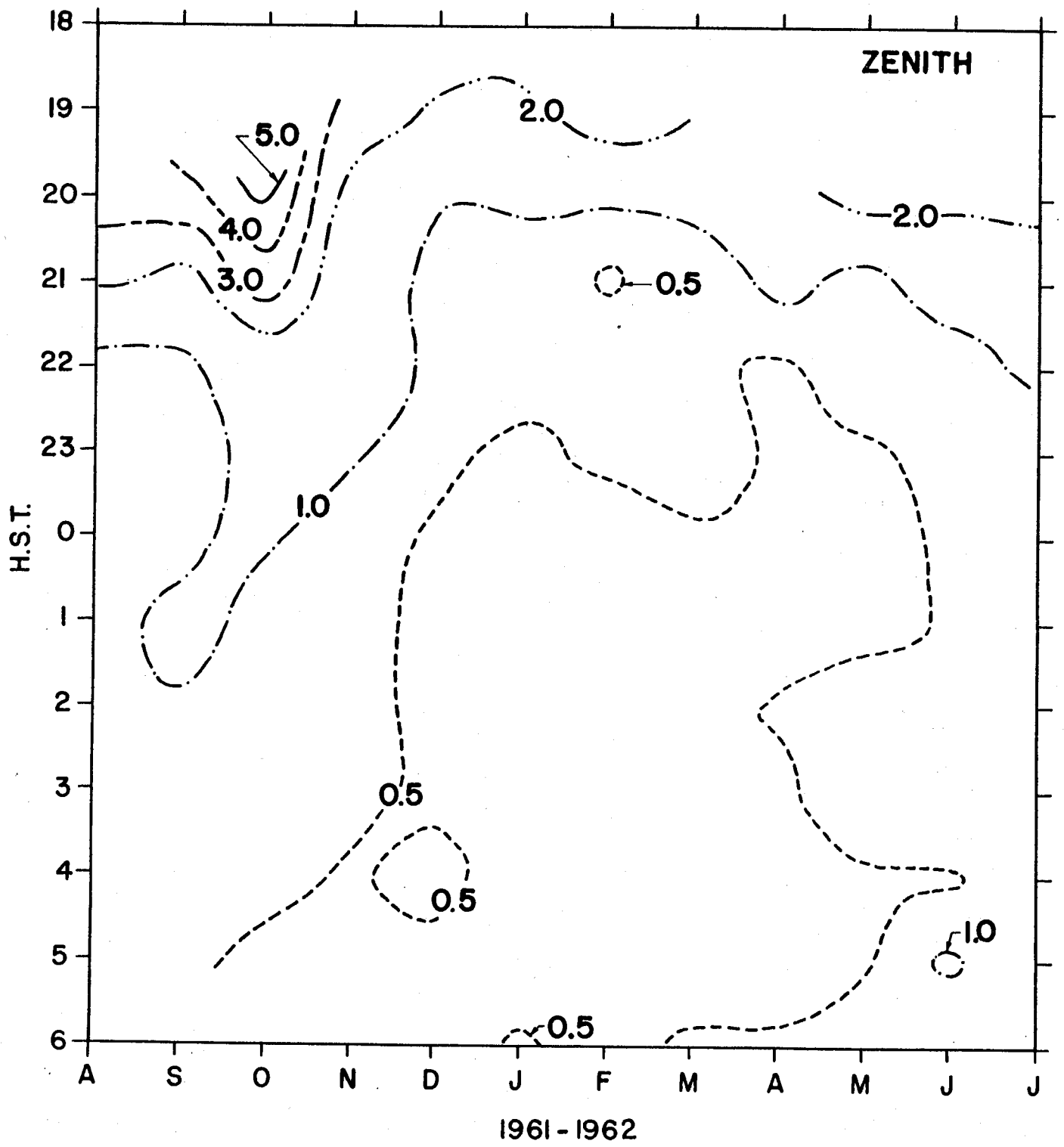


Figure 32

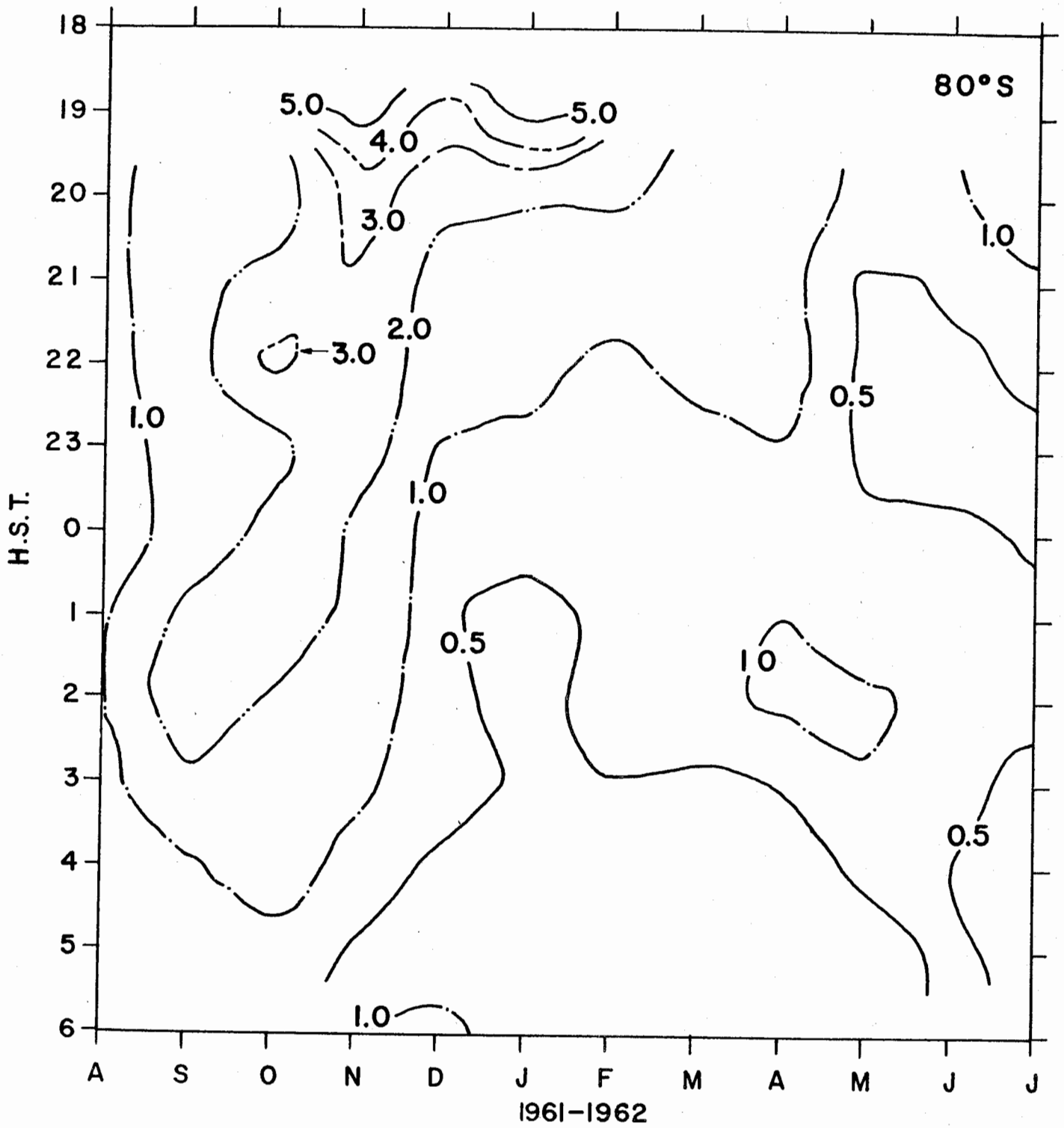


Figure 33

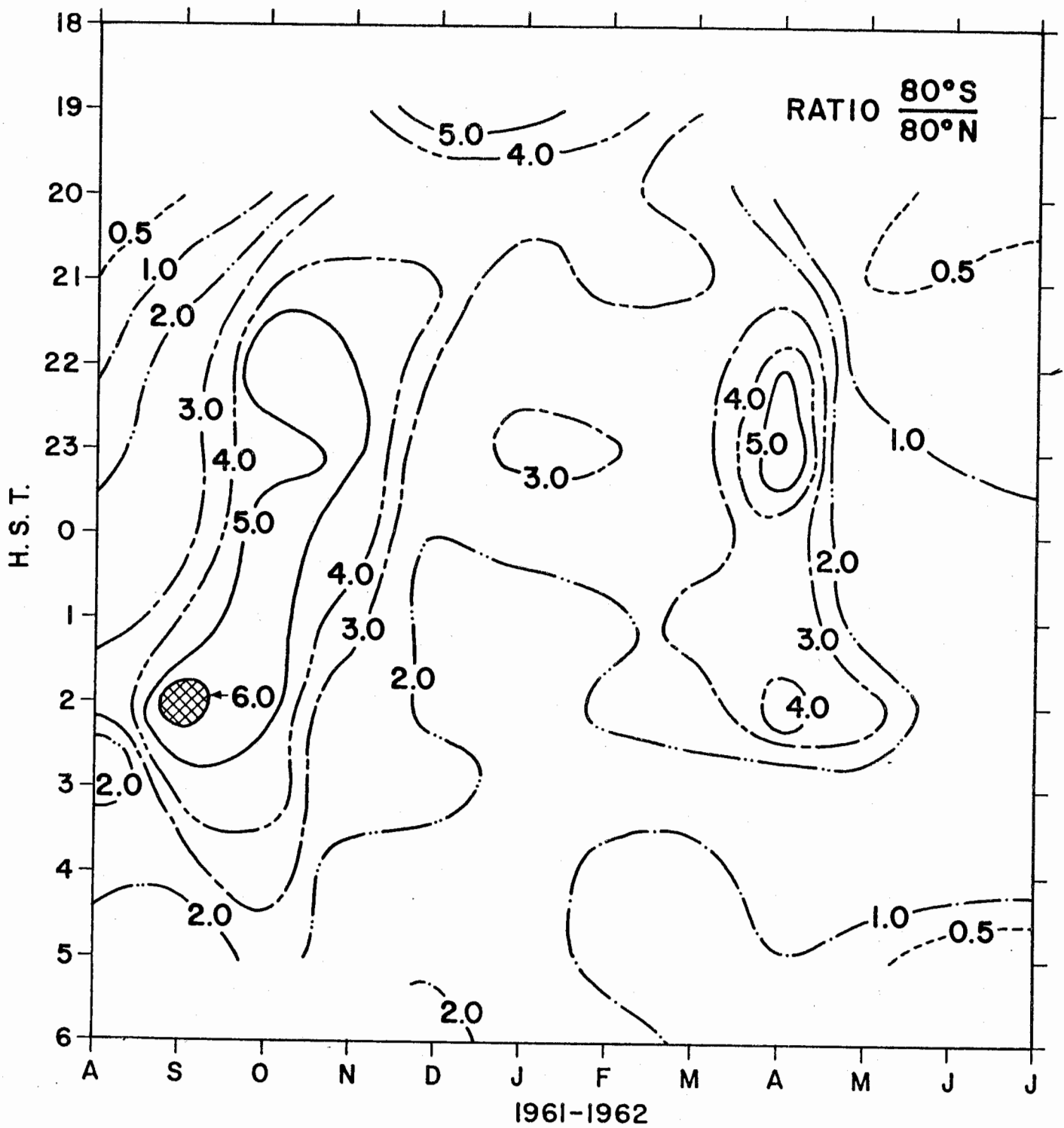


Figure 34

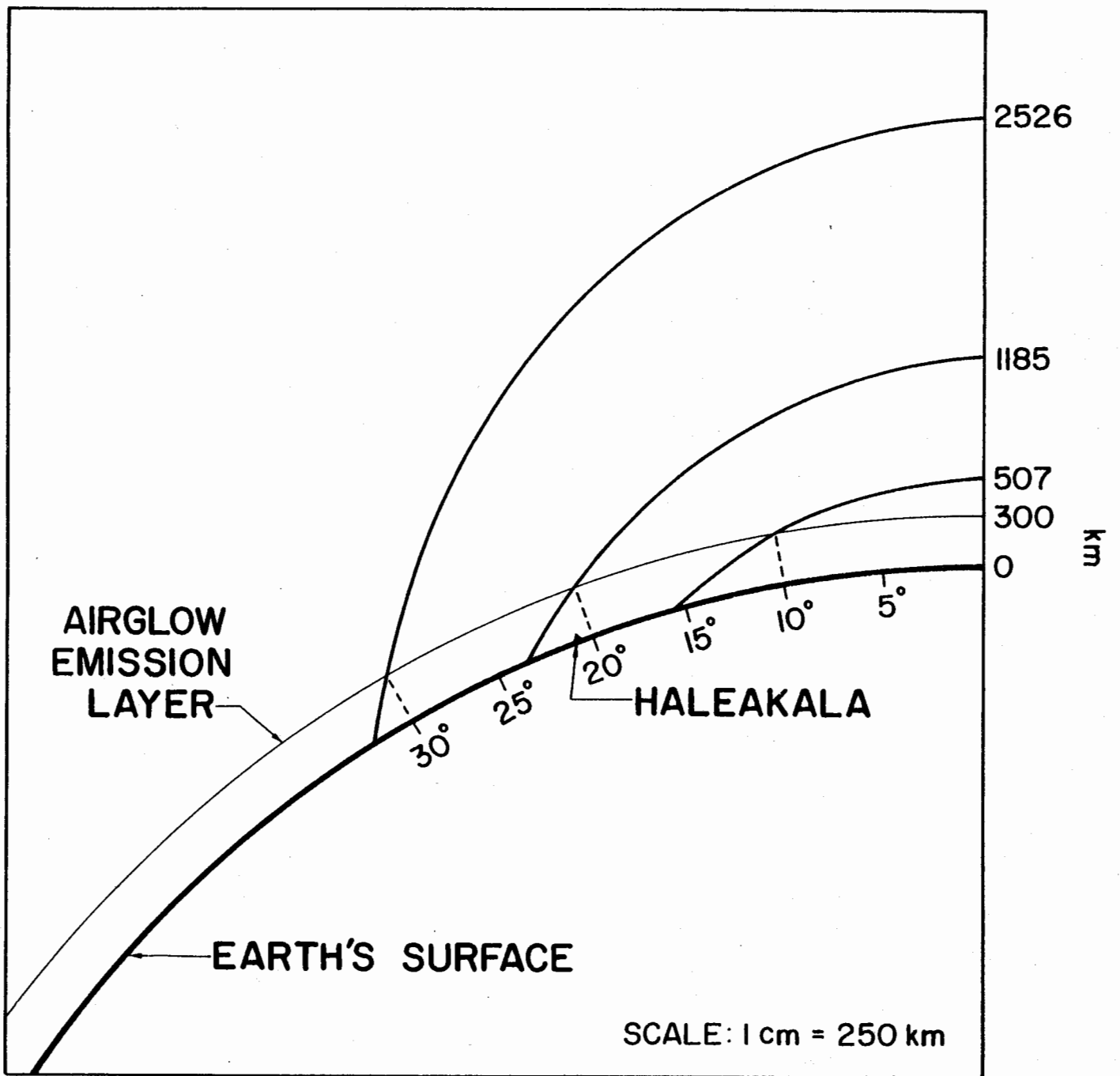


Figure 35

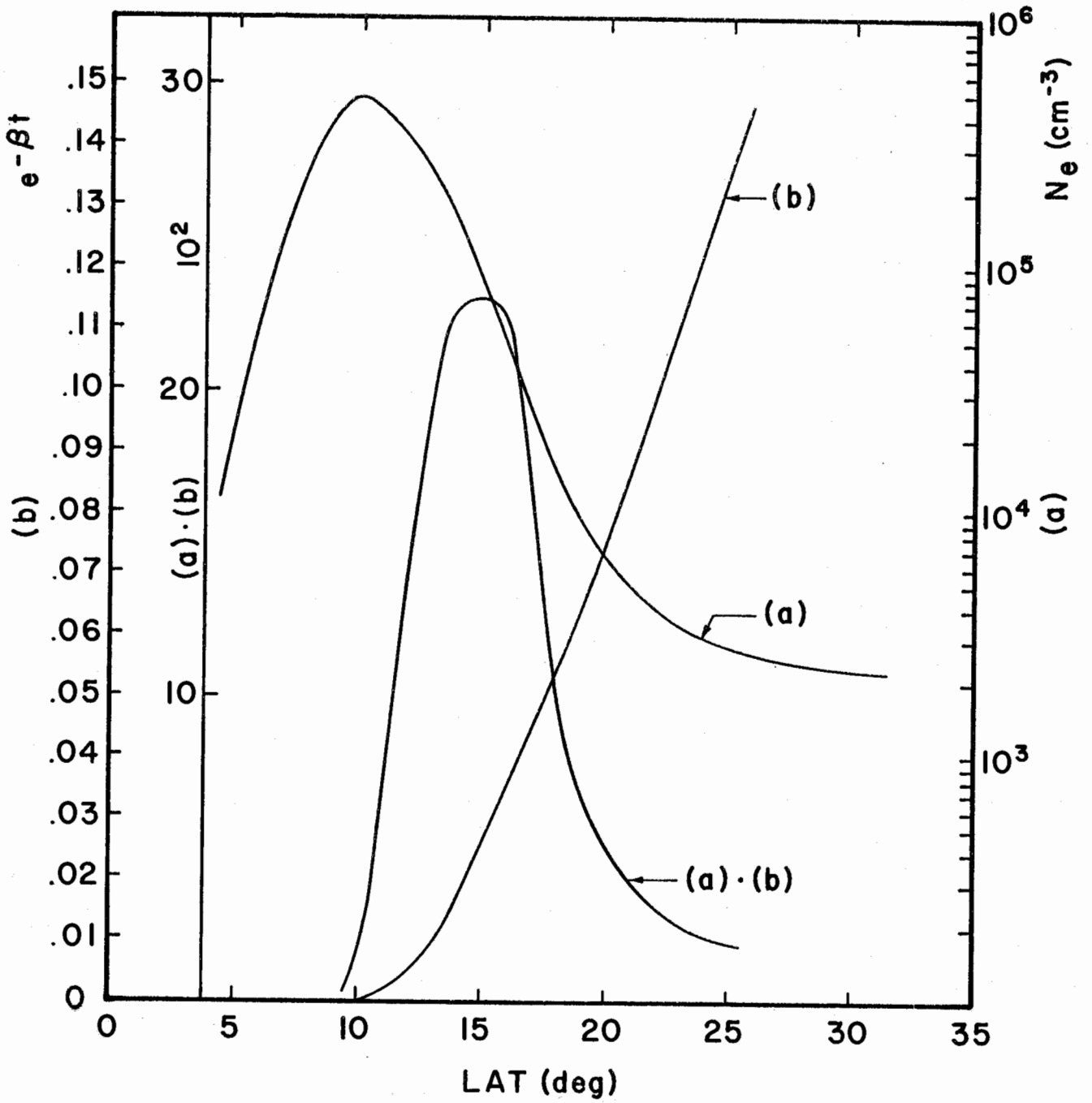


Figure 36